

1 **3D unified geological model of the Milk River Transboundary Aquifer**

2 **(Alberta, Canada-Montana, USA)**

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26 **Abstract**

27  
28 The Milk River Transboundary Aquifer (Canada/USA) has been so intensively used over the  
29 twentieth century that concerns have risen about the durability of this resource since the mid-  
30 1950s. This aquifer actually corresponds to the middle Virgelle Member of the Upper Cretaceous  
31 Milk River Formation (called Eagle Formation in Montana). To assess the conditions needed for  
32 a sustainable use of the aquifer, a comprehensive and unified portrait of the aquifer is needed  
33 across its international boundary. The stratigraphic framework and geometry of geological units  
34 on both sides of the international border were thus unified in a 50 000 km<sup>2</sup> 3D geological model.  
35 The Virgelle Member is 0 to 60 m thick and it subcrops near the border and along both sides of  
36 the Sweetgrass Arch. It dips away from the subcrop areas in a semi-radial pattern. The Medicine  
37 Hat gas field hosted by the Alderson Member (Alberta), which is separated from the other  
38 members by a regional unconformity, and the Tiger Ridge gas field near the Bears Paw  
39 Mountains (Montana) limit the extent of the aquifer. The unified 3D geological model forms the  
40 necessary basis for conceptual and numerical hydrogeological models of the Milk River Aquifer.

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42 **Résumé**

43  
44 L'aquifère transfrontalier Milk River (Canada/USA) a été intensivement sollicité pendant le 20<sup>ème</sup>  
45 siècle, si bien que des inquiétudes concernant la durabilité de cette ressource sont apparues dès le  
46 milieu des années 1950. Cet aquifère correspond en réalité au Membre Virgelle de la Formation  
47 Milk River (appelé Formation Eagle au Montana) datant du Crétacé Supérieur. Pour évaluer les  
48 conditions nécessaires à un usage durable de l'aquifère, un portrait complet et unifié de l'aquifère  
49 est nécessaire à travers sa frontière internationale. Le cadre stratigraphique et la géométrie des

50 unités géologiques des deux côtés de la frontière internationale ont été ainsi unifiés dans un  
51 modèle géologique 3D de 50 000 km<sup>2</sup>. Le Membre Virgelle a une épaisseur allant de 0 à 60  
52 mètres et sous-affleure près de la frontière et le long des deux côtés de l'arche Sweetgrass. Il  
53 plonge depuis les zones de sous-affleurement en suivant une disposition semi-radiale. Le champ  
54 de gaz Medicine Hat, contenu par le Membre Alderson (Alberta) qui est séparé de la Formation  
55 Milk River par une discordance régionale et le champ de gaz Tiger Ridge près des Montagnes  
56 Bears Paw (Montana) délimitent l'aquifère. Ce modèle géologique 3D unifié forme la base  
57 indispensable aux modèles hydrogéologiques conceptuel et numérique de l'aquifère Milk River.

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59 **Keywords:** transboundary aquifer, geological model, hydrostratigraphy, Canada, USA

60

## 61 **Introduction**

62

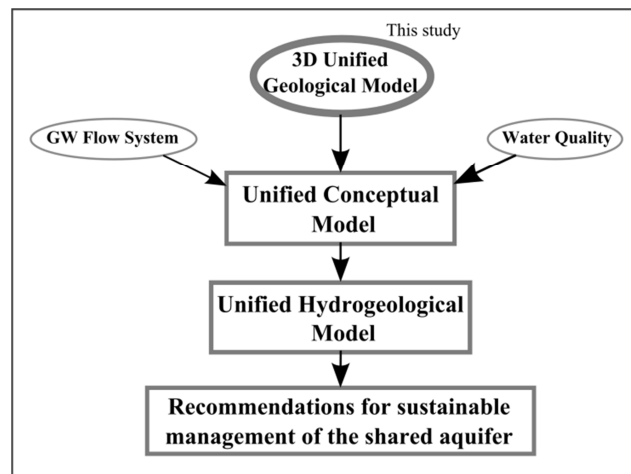
63 The Milk River Aquifer straddles southern Alberta (Canada) and northern Montana (USA) in a  
64 water-short semi-arid region (Government of Alberta, Alberta water for life 2006). This confined  
65 sandstone aquifer corresponds more specifically to the Virgelle Member of the Milk River  
66 Formation (Eagle Formation in Montana). The extensive use of this resource over the 20<sup>th</sup>  
67 century has led to a major drop in water levels locally, and concerns about the durability of the  
68 resources have been raised since the mid-1950s (Meyboom 1960; Borneuf 1976). More recently,  
69 AGRA (1998) published a depletion study of the aquifer and made recommendations for the  
70 conservation of this resource. A 5-year conservation program in southern Alberta followed this  
71 study; however the Milk River Aquifer is still solicited on both sides of the international border  
72 in the absence of an agreement between the USA and Canada on the use of this shared resource.

73

74 The Milk River Aquifer has been the object of many studies throughout the 20<sup>th</sup> century; however  
75 most of them were limited by the Canada/USA border (Meyboom 1960; Hendry et al. 1991;  
76 Alberta Innovates Technology Future 2010), thus preventing a full understanding of the aquifer  
77 dynamics. Since geological formations in Alberta and Montana are not defined or named the  
78 same way, several stratigraphic charts of the Milk River Formation subdivision have evolved  
79 considerably in each country, making transboundary studies challenging. An integrated portrait of  
80 the aquifer is, however, necessary to assess the conditions needed for its sustainable shared use.

81  
82 The objective of this study is to overcome transboundary limitations by providing unified  
83 geological and conceptual hydrogeological models of the Milk River Aquifer. The term “unified”  
84 means that the study follows the natural limits of the aquifer, not interrupted by the international  
85 border. A consistent nomenclature is required to describe the stratigraphic and hydrostratigraphic  
86 units across the international boundary. A first component of the conceptual model is the 3D  
87 unified geological model of the aquifer presented in this paper (Fig. 1). For this purpose,  
88 geological data on both sides of the international border were gathered and assembled  
89 consistently in the light of the latest correlation and geological work (Payenberg 2002b;  
90 Payenberg et al. 2002; O’Connell 2014). The 3D geological model thus provides a common  
91 stratigraphic framework for hydrogeological applications, as done by Ross et al. (2005) and it  
92 represents a prerequisite for a representative transboundary hydrogeological numerical model,  
93 which in turn will lead to recommendations for the sound management of this shared resource.

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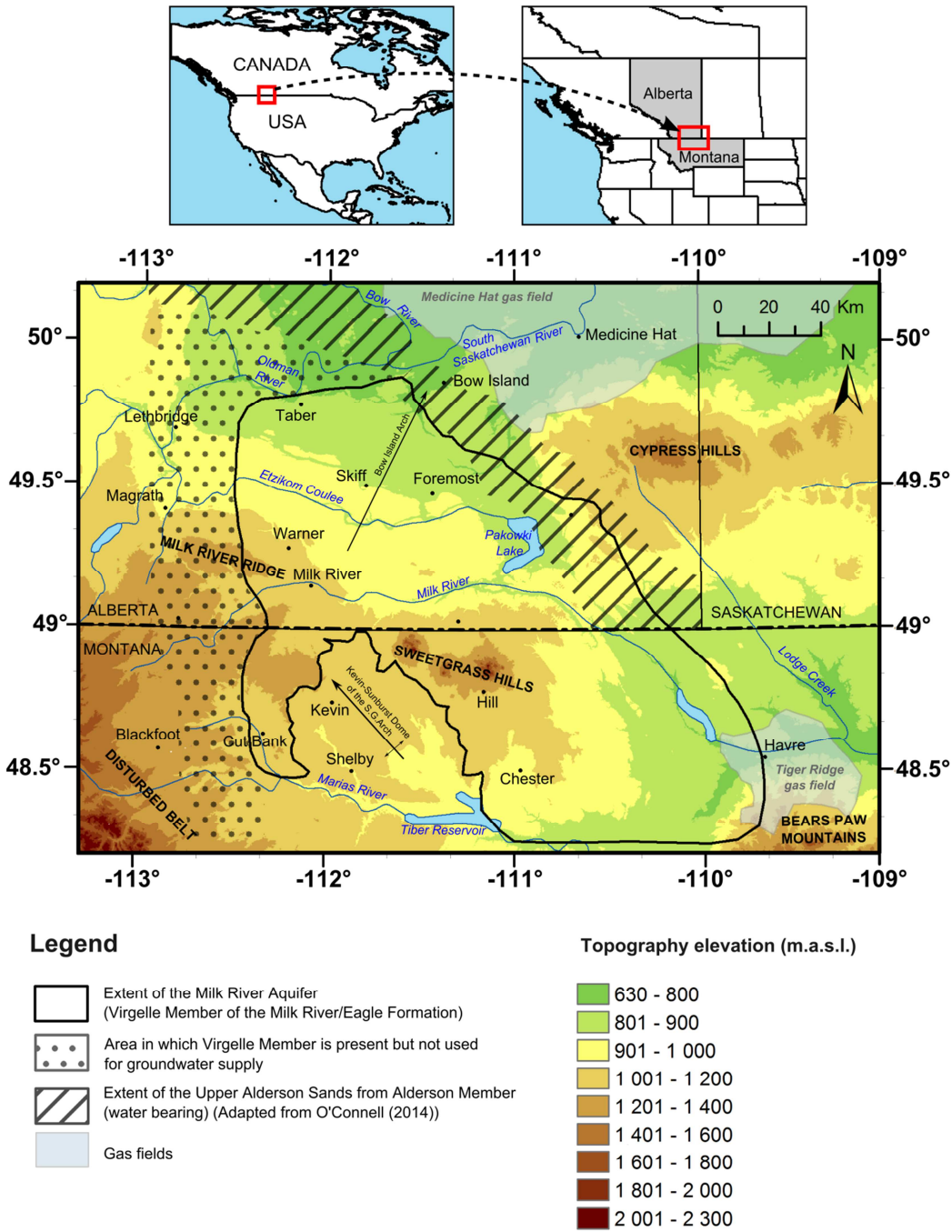
95  
 96 **Fig. 1.** Successive stages of the Milk River transboundary aquifer study; this paper presents the  
 97 initial step, the development of a 3D unified geological model.

98  
 99 **Study area**

100  
 101 The study area extends over about 50 000 km<sup>2</sup> in southern Alberta (Canada) and northern  
 102 Montana (USA) (Fig. 2). The study area ranges from longitude -110.0° to -113.0° and from  
 103 latitude 48.2° to 50.3°. The study area is bounded by the edge of the Disturbed Belt (indicated on  
 104 Fig. 2), on the west, and reaches the Saskatchewan border and the Bears Paw Mountains, on the  
 105 east. The southern limit is north of the Marias River in Montana. The northern limit is the  
 106 southern edge of the Medicine Hat gas field in Alberta.

107  
 108 The main structural feature in the study area is the Sweetgrass Arch, composed of the Kevin-  
 109 Sunburst Dome, the Bow Island Arch and the Sweetgrass Hills. The Sweetgrass Arch in Montana  
 110 is composed of the Kevin-Sunburst Dome and the South Arch. The South Arch is outside the  
 111 study area. The Sweetgrass Hills are on the eastern flank of the Sweetgrass Arch. The Sweetgrass

112 Hills are an ensemble of three buttes (2 100 m altitude) near the Canada/US border. The limits of  
 113 the model are explained in the section “Geological and Hydrostratigraphic Settings”.



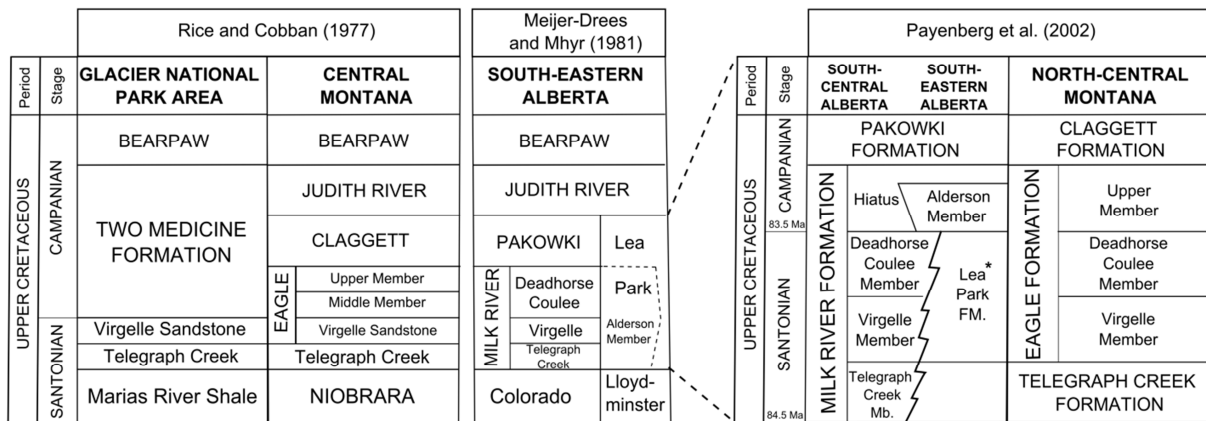
114  
 115 **Fig. 2.** Study area and extent of the Milk River Aquifer. “S.G. Arch” stands for Sweet Grass  
 116 Arch.

117

118 **Correlations**

119

120 Prior to discussing the geological and hydrogeological contexts of the Milk River Aquifer, it is  
 121 necessary to establish the correlations between geological units across the international boundary  
 122 and define the nomenclature that will be used in the remainder of this paper. As the  
 123 characterization of the Upper Cretaceous Milk River Formation (Eagle Formation in Montana)  
 124 progressed, the stratigraphic nomenclature evolved significantly during the twentieth century  
 125 (Fig. 3). The stratigraphic charts not only differ between southern Alberta and northern Montana  
 126 but also within northern Montana (east and west of the Sweetgrass Arch). As early as 1917,  
 127 Stebinger (1917b) described the differences between the geologic sections east and west of the  
 128 112<sup>th</sup> meridian in Montana.



129

130 **Fig. 3.** Comparative stratigraphic nomenclatures existing in the study area (modified from Rice  
 131 and Cobban 1977; Meijer-Drees and Mhyr 1981 and Payenberg et al. 2002).\* The Lea Park  
 132 Formation may be drawn all the way through the Pakowki Formation, following Meijer-Drees  
 133 and Mhyr (1981) and Dawson et al. (1994).

134

135  
136 The Milk River Formation of southern Alberta was first identified by Dowling (1915, 1917) as  
137 the “Milk River Sandstone” in his study of the Southern Plains of Alberta. The Milk River  
138 Sandstone consisted then of two parts (lower and upper) bearing successive terminologies (Evans  
139 1931; Russell and Landes 1940). A three-part subdivision of the Milk River Formation was  
140 introduced by Tovell (1956), including the Transition beds, Virgelle and Deadhorse Coulee  
141 (Tovell 1956, cited by Meijer-Drees and Mhyr 1981). In Montana, the upper and lower parts of  
142 the Milk River Sandstone were referred to as Upper Eagle and Virgelle by Williams and Dyer  
143 (1930). Meyboom (1960) also equated the lower part of the Milk River Formation to the Virgelle  
144 Member of Eagle Sandstone in Montana.

145  
146 In Montana, Stanton et al. (1905) first described the stratigraphy of Upper Cretaceous rocks in  
147 northern and central Montana and in Canada. They defined the Eagle Formation (named by Weed  
148 1899 from Eagle Creek, a tributary of the Missouri River) as massive white sandstone overlain by  
149 softer beds consisting of alternating sandstone, shale, and many beds and seams of lignite. They  
150 also noticed that small black pebbles occurred at the top. Rice (1980) divided the Eagle  
151 Formation into three members: the basal Virgelle Member, and the unnamed middle and upper  
152 members. Stanton et al. (1905) established that the overlying Pakowki Shale in Alberta was  
153 identical to the Claggett Shale in Montana. They also showed that the Belly River Group that  
154 overlies the Pakowki Shale in Alberta was identical to the Judith River in Montana. Payenberg et  
155 al. (2002) confirmed that Claggett and Pakowki are of the same age. In the northeastern part of  
156 the study area, the stratigraphic equivalent of the Pakowki Formation is the Lea Park Formation  
157 (Williams and Dyer 1930).

158  
8



159 Meijer-Drees and Mhyr (1981) proposed a stratigraphic nomenclature for southeastern Alberta.  
160 They defined the Milk River Formation as the stratigraphic equivalent of the Telegraph Creek  
161 Formation and Eagle Sandstone defined by Rice and Cobban (1977). They also defined the  
162 Deadhorse Coulee Member in Southeastern Alberta as occupying the same interval as the upper  
163 and middle members of the Eagle Sandstone. In the Sweetgrass Hills area, Tuck (1993) named  
164 the interval between Virgelle and Claggett the “upper part” of Eagle sandstone. This upper part  
165 consisted of interbedded shale, siltstone, sandstone, and coal, and bears close resemblance to the  
166 Deadhorse Coulee Member.

167  
168 Until recently, a clear regional correlation of the Milk River Formation and the Eagle Formation  
169 was not possible, due to differences in lithology and time-range (Russell 1970) and the limited  
170 and remote exposure of the Eagle Formation within northern Montana (Payenberg et al. 2003).  
171 Russell (1970) revealed misunderstandings made in previous stratigraphic correlations. Thirty  
172 years after Russell’s (1970) attempt at correlating between the Milk River and Eagle Formations,  
173 Payenberg (2002b) and Payenberg et al. (2002) reevaluated the lithostratigraphic and  
174 chronostratigraphic relationships of Alberta and Montana Upper Cretaceous rocks. They used  
175 recent advances in geochronology, magnetostratigraphy, and a paleontological database. The  
176 work of Payenberg et al. (2002) provided a clearer litho- and chrono-stratigraphic framework of  
177 the study area; it particularly showed that the Telegraph Creek, Virgelle, and Deadhorse Coulee  
178 Members are continuous and correlative across the international border (Payenberg 2002a), and  
179 also introduced the Alderson Member of the Lea Park Formation in the southern Alberta  
180 nomenclature (Fig. 3).

181

182 The proposed nomenclature in the present study (Fig. 4) is based on the previous works of  
 183 Meijer-Drees and Mhyr (1981), Payenberg et al. (2002) and Rice and Cobban (1977). The study  
 184 area is divided in four zones (Fig. 5a), each with a distinct succession of geological units. These  
 185 zones are defined as follows:

186 Zone 1: South-western part of the study area in Alberta, southwest of the Virgelle depositional  
 187 limit;

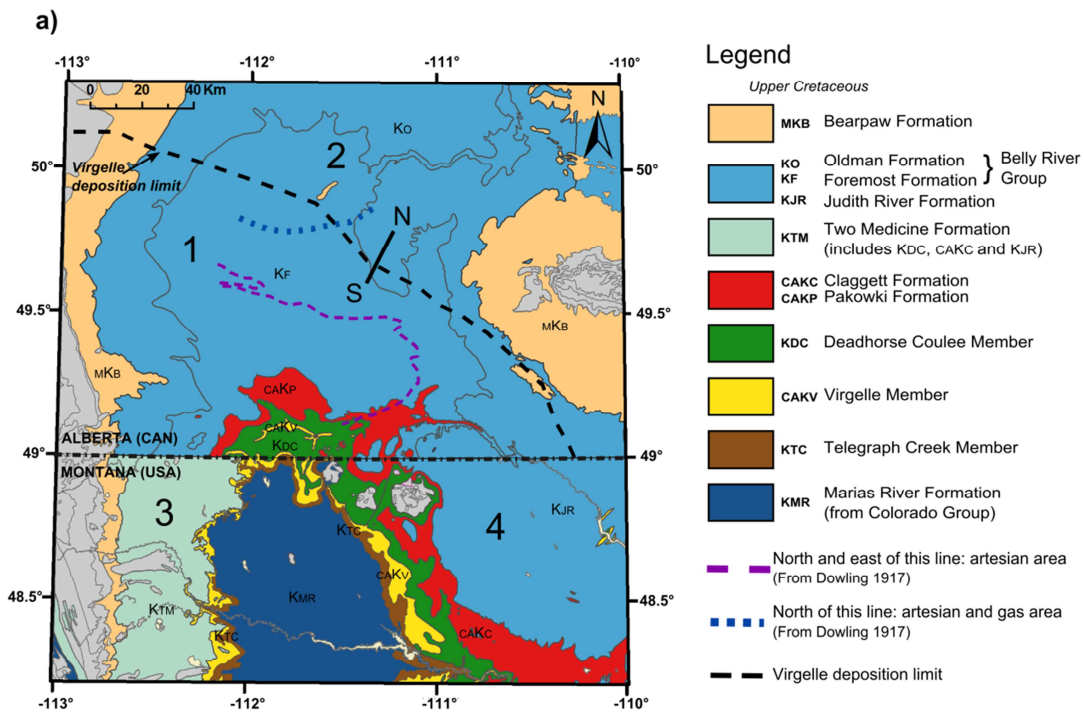
188 Zone 2: North-eastern part of the study area in Alberta, northeast of Virgelle depositional limit;

189 Zone 3: North-western Montana, west of the Sweetgrass Arch;

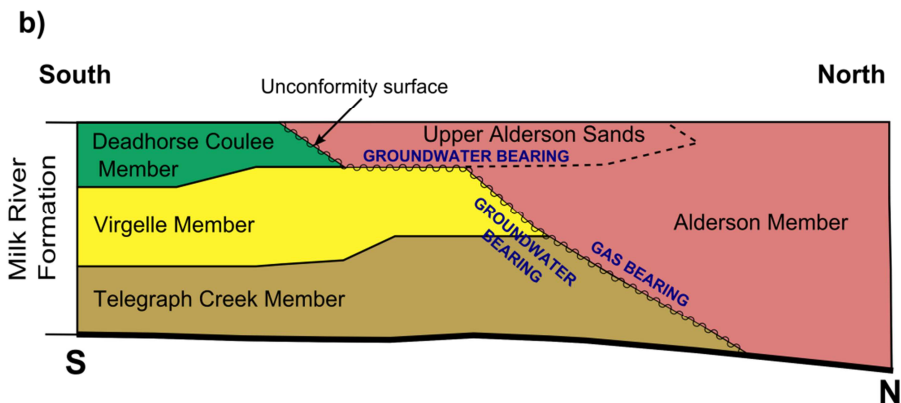
190 Zone 4: Northern Montana, east of the Sweetgrass Arch.

|                      | Zone 1               | Zone 2                         | Zone 3              | Zone 4               |
|----------------------|----------------------|--------------------------------|---------------------|----------------------|
|                      | Bearpaw Fm           | Bearpaw Fm                     | Bearpaw Fm          | Bearpaw Fm           |
|                      | Belly River Group    | Belly River Group              | Two                 | Judith River Fm      |
|                      | Pakowki Fm           | Pakowki Fm                     | Medicine            | Claggett Fm          |
| Milk River Formation | Deadhorse Coulee Mbr | Lea Park Fm<br>Alderson Member | Fm                  | Deadhorse Coulee Mbr |
|                      | Virgelle Mbr         |                                | Virgelle Mbr        | Virgelle Mbr         |
|                      | Telegraph Creek Mbr  |                                | Telegraph Creek Mbr | Telegraph Creek Fm   |
|                      | Colorado Group       | Colorado Group                 | Colorado Group      | Colorado Group       |

191  
 192 **Fig. 4.** Proposed stratigraphic nomenclature in the present study and representation of the  
 193 geological formations described in the 3D geological model (zones shown on Fig. 5a). Note that  
 194 the Two Medicine Formation has been subdivided into three members (Judith River Formation,  
 195 Claggett Formation and Deadhorse Coulee Member equivalents). The Bearpaw Formation  
 196 box also includes the surficial sediments in the 3D model.



198



199

200 **Fig. 5. a)** Bedrock geological map of the study area (Adapted from Okulitch et al. 1996). **b)**

201 Cross-section (indicated by “S-N” in Fig. 5a) showing the unconformity surface separating the

202 Alderson Member from the three other members of the Milk River Formation. The encasing units

203 are not represented on this cross section (Adapted from O’Connell 2014).

204

205

206 **Geological and Hydrostratigraphic Settings**

207

208 The geology of the study area can be described as a succession of marine and continental  
 209 sediments that were deposited as the Upper Cretaceous Interior Sea level fluctuated (Russell  
 210 1970). The Upper Cretaceous strata are briefly described below and represented on the bedrock  
 211 geological map in Fig. 5a. Their hydrostratigraphic role is indicated in Fig. 6 and briefly  
 212 described in the present section.

| Period                  | STRATIGRAPHY         | HYDROSTRATIGRAPHY       |                     |
|-------------------------|----------------------|-------------------------|---------------------|
| <b>UPPER CRETACEOUS</b> | Bearpaw Formation    | Bearpaw Aquitard        |                     |
|                         | Belly River Group    | Dinosaur Park Formation | Belly River Aquifer |
|                         |                      | Oldman Formation        |                     |
|                         |                      | Foremost Formation      |                     |
|                         | Pakowki Formation    | Pakowki Aquitard        |                     |
|                         | Milk River Formation | Milk River Aquifer      |                     |
|                         | Colorado Group       | Colorado Aquitard       |                     |

213

214 **Fig. 6.** Stratigraphy and hydrostratigraphy of the main geological units of the study area.

215

216

217 **Colorado Group**

218 The Colorado Group (middle Albian to Santonian) underlies the entire study area (Fig. 4). It was  
 219 deposited during marine conditions in a moderate-depth sea environment. The Colorado Group  
 220 consists mainly of dark grey to black bentonitic marine shale. It ranges in thickness from 500 to

221 600 m in southern Alberta and from 450 to 500 m in north central Montana (Hendry et al. 1991;  
222 Cobban et al. 1976). The upper boundary of the Colorado Group is commonly taken at the First  
223 White Speckled Shales (Meyboom 1960). The Colorado Group is not exposed in southern  
224 Alberta (Dyer and Williams 1930) but it outcrops widely over 5 counties in northern Montana,  
225 from the Sweetgrass Hills to Great Falls (Cobban et al. 1976). It constitutes a regional aquitard in  
226 the study area (Fig. 6) with a hydraulic conductivity ranging from  $10^{-14}$  to  $10^{-10}$  m/s (Hendry and  
227 Schwartz 1988). The Colorado Group is overlain by the Milk River Formation and thus the  
228 interface between these units constitutes the datum plane of the present regional study.

229

### 230 **Milk River/Eagle Formation**

231 The Milk River Formation (Eagle Sandstone in Montana) is a regressive clastic wedge deposited  
232 during the Late Cretaceous (Rice 1980; Payenberg et al. 2001). The Milk River Formation has  
233 been traditionally subdivided into three members: the basal Telegraph Creek Member, the middle  
234 Virgelle Member and the upper Deadhorse Coulee Member.

235 The Milk River Formation is 150 m thick in the southwest corner of the Canadian part of the  
236 study area and thins towards the northeast (O'Connell 2014). It subcrops in an area of 14  
237 townships in southern Alberta near the border, in rings around the Sweetgrass Hills and along  
238 both sides of the Sweetgrass Arch. According to Russell (1970): "There is little difference  
239 between the western and eastern developments." The Milk River Formation dips gently to the  
240 north, east and west, from the subcrop areas following a radial or "fan-like" pattern (Meyboom  
241 1960; Schwartz and Muehlenbachs 1979; Toth and Corbet 1986). The Milk River Formation is  
242 confined below and above by the low-permeability shales of the Colorado Group and  
243 Pakowki/Claggett Formations, respectively.

244

245 The Milk River Aquifer is within the Milk River Formation. The middle Virgelle Member is the  
246 most important aquifer within the formation. The Milk River Aquifer is a confined and inclined  
247 aquifer, which locally shows flowing artesian conditions. About 200 flowing artesian wells were  
248 inventoried in the 1960s in southern Alberta (Meyboom, 1960). In Montana, flowing artesian  
249 water occurs in much of the area of Cut Bank, Montana due to the westward dip of the formation  
250 (Zimmerman 1967). Tuck (1993) also highlighted some flowing artesian wells in the Sweetgrass  
251 Hills area. However, many wells have lost their artesian flow because of the intensive use of this  
252 resource. Nowadays, the flowing artesian areas are located in the vicinity of Pakowki Lake and  
253 north of the study area, which is still consistent with the flowing artesian limit drawn by Dowling  
254 (1917) (Fig. 5a).

255  
256 Meyboom (1960) showed that the recharge areas of the aquifer were located mainly in the  
257 concentric outcrops around the Sweetgrass Hills and to a lesser extent at the subcrop area near the  
258 international border. The main discharge areas are located at the pumping or flowing wells of the  
259 study area, with small natural discharge (Meyboom 1960).

260  
261 Groundwater flow in the Milk River Aquifer follows the regional dip of the Milk River/Eagle  
262 Formation. In Alberta the general flow is semi-radial from the topographic highs of the  
263 Sweetgrass Hills to the north, west and east (Hendry and Schwartz, 1988). In Montana,  
264 groundwater flows from the Sweetgrass Hills to the east and south-east and also from the subcrop  
265 areas west of the Sweetgrass arch to the west and south (Zimmerman 1967; Levings 1982).  
266 Therefore, there are two transboundary flow paths in the study area: 1) from the Sweetgrass Hills  
267 to the north, and 2) from north of the Cut Bank, Montana, area to the north (Zimmerman 1967;  
268 Tuck 1993).

269  
270 In southern Alberta, the Alderson Member of Lea Park Formation (Meijer-Drees and Mhyr  
271 (1981) unconformably overlays the Milk River Formation in the north, northeast and east of the  
272 study area (Fig. 5b). The Alderson Member is a lateral sandy shale equivalent to the Milk River  
273 Formation and some authors (Hendry et al 1991) described the relation between the units as a  
274 facies change. The Alderson Member has also been included in the Milk River Formation as a  
275 fourth member (Payenberg et al. 2002; O'Connell 2014). However, it is much younger than the  
276 other three members, is separated from them by a regional unconformity surface and is not  
277 present in Montana (Payenberg et al. 2003; O'Connell 2014). In the framework of the present  
278 study, the decision was made to include the Alderson Member as a member of the Lea Park  
279 Formation as also done by other authors (e.g. Meijer-Drees and Mhyr 1981). The Alderson  
280 Member is gas-bearing; it contains the Milk River gas field (or Medicine Hat gas field, Fig. 2), a  
281 natural limit of the 3D geological model. Another natural limit is imposed by the Tiger Ridge gas  
282 field in Montana, hosted by the Eagle Sandstone in the Bears Paw Mountains area (Fig. 2)  
283 (Gautier and Rice 1982), and constitutes another natural boundary for the present 3D geological  
284 model.

285  
286 **Telegraph Creek Member/Formation**

287 The Telegraph Creek Member is a transitional unit between the shale of the Colorado Group and  
288 the massive sandstone of the Virgelle Member of the Milk River Formation. It is interpreted as  
289 deposits of an offshore to shore-face transition (Payenberg 2002a). The Telegraph Creek Member  
290 consists of thinly interbedded sandy shale, siltstone and fine-grained shaly sandstone. The  
291 Telegraph Creek is 36 to 52 m thick in the Cut Bank, Montana, area where it has formation status  
292 (Cobban 1950; Payenberg et al. 2001) and it is 30 to 52 m thick near the Sweetgrass Hills

293 (Zimmerman 1967; Tuck 1993). The Telegraph Creek transition zone was originally included in  
294 the Virgelle Sandstone in north-central and northwestern Montana by Stebinger (1915, 1917a).

295

## 296 **Virgelle Member**

297 The Virgelle Member gradationally overlies the Telegraph Creek Member (Meijer-Drees and  
298 Mhyr 1981). It consists of grey to buff, thick bedded, fine to medium grained sandstone with  
299 thinly bedded siltstone (Tuck 1993). The Virgelle Member was deposited during a regression  
300 sequence and is interpreted as a shore-face to foreshore sandstone (Rice 1980). It is up to 69 m  
301 thick in southern Alberta and varies from 15 to 60 m thick on the west side of the Sweetgrass  
302 Arch (Lorenz 1981; O'Connell 2014). The Virgelle Member is not present in southwestern  
303 Saskatchewan or central Alberta because it is truncated by the regional unconformity surface  
304 separating the Milk River Formation and the Alderson Member. The Virgelle outcrops along the  
305 Milk River in southern Alberta over approximately 25 km in Township 1 and 2, Ranges 12 to 15  
306 (Meyboom 1960). It also outcrops on both sides of the Sweetgrass Arch, in continuous and  
307 narrow belts (Fig. 5a).

308

309 The Virgelle Member massive sandstone is the most important aquifer part of the Milk River  
310 Formation and therefore constitutes the Milk River Aquifer. The hydraulic conductivity of the  
311 Virgelle sandstone in southern Alberta ranges from  $10^{-8}$  to  $10^{-6}$  m/s (Persram 1992 unpublished,  
312 cited by AGRA 1998). South-east of the town of Milk River, the hydraulic conductivity of the  
313 Virgelle is  $1.8 \times 10^{-7}$  m/s (Robertson 1988). The limits of the Milk River Aquifer are shown on  
314 Fig. 2. They correspond to the area in which the Virgelle Member exists and is exploited.  
315 However the 3D geological model extends farther west, to longitude  $-113^{\circ}$ , where the Virgelle  
316 Member continues in the subsurface but is too deep ( $> 400$  m) to be used for groundwater supply



317 (Stantec 2002). In the framework of the present 3D geological model, the internal stratigraphy,  
318 lithofacies, and depositional environment of the Virgelle Member are not further discussed.  
319 However, these aspects were studied by Rice (1976; 1980) in northern Montana and Meyer  
320 (1998) as well as Meyer and Krause (2006) in southern Alberta.

321

### 322 **Deadhorse Coulee Member**

323 The Deadhorse Coulee Member (DHC) represents the upper part of the Milk River Formation as  
324 named by Tovell (1956). It is a non-marine unit deposited in the coastal plain environments  
325 landward of the Virgelle shore-faces (O'Connell 2014). This well-defined unit consists  
326 predominantly of interbedded shale, siltstone and fine grained sandstone with coal seams  
327 (Payenberg 2002). The Deadhorse Coulee has a maximum thickness of 60 m in southern Alberta  
328 and thins northeastward to approximately 10 m east of the zero edge (from T1 R5 W4 to T13 R22  
329 W4) (O'Connell 2014). In northern Montana, the Deadhorse Coulee equivalent is the unnamed  
330 middle member of Eagle Formation (Payenberg et al. 2001). The contact between Deadhorse  
331 Coulee and the overlying Pakowki /Claggett Formation is marked by a thin (but laterally  
332 continuous) bed of dark grey to black polished chert pebbles, which is interpreted as a  
333 transgressive lag overlying a regional unconformity surface (Russell 1970; O'Connell 2014). The  
334 Deadhorse Coulee Member constitutes a low permeability unit that overlies the Virgelle Member  
335 (O'Connell 2014).

336

### 337 **Alderson Member**

338 The Alderson Member was originally considered the lower member of the Lea Park Formation  
339 and a stratigraphic equivalent to the Milk River Formation (Meijer-Drees and Mhyr 1981)(Fig.  
340 3). It was entirely deposited in proximal to distal offshore marine environments (O'Connell

341 2014). In southern Alberta, the Alderson Member is present just northeast of the depositional  
342 limit of the Virgelle sandstone (Meijer-Drees and Mhyr 1981). The lithology of the Alderson  
343 Member consists of interbedded very fine-grained sand, silt and mud (O’Connell 2014). The sand  
344 content increases in the upper part (Meijer-Drees and Mhyr 1981). O’Connell (2014) includes the  
345 Alderson Member as the youngest member of the Milk River Formation which is 100 m thick in  
346 the northeast corner of the study area. The Alderson Member is younger than the Telegraph  
347 Creek, Virgelle and Deadhorse Coulee Members of the Milk River Formation and is separated  
348 from them by a regional unconformity representing a large time-gap (O’Connell 2014; Payenberg  
349 2003). It therefore physically overlies and overlaps the erosional edges of the other three  
350 members. The Alderson Member is about 85 m thick in southeastern Alberta (Meijer-Drees and  
351 Mhyr 1981).

352  
353 The Alderson Member hosts the Medicine Hat gas field (Hamblin and Lee 1997; Fig. 2).  
354 However, the upper part of the Alderson Member contains two distinct large sand bodies which  
355 form a regional aquifer in southern Alberta. This unit is named the Upper Alderson Sands by  
356 O’Connell (2014). It covers an area of 74 townships and has a NW-SE trend (Fig. 2). The Upper  
357 Alderson Sands forms small lobate sand bodies. According to O’Connell (2014): “The Virgelle  
358 and Upper Alderson aquifers are separated from each other by muddy sediments of the Alderson  
359 and Deadhorse Coulee members, but they are locally in contact at the Virgelle erosional edge and  
360 water flow between the two aquifers is likely.”

361  
362 **Pakowki Formation/ Claggett Shale**  
363 The Milk River Formation is overlain by a thick unit of marine shales, the Pakowki Formation  
364 (Claggett Shale equivalent in Montana). The Pakowki and Claggett Formations consist of thinly

365 bedded, black marine shales, with few sandstone beds (Tovell 1956 cited by Payenberg et al.  
366 2003). The Pakowki Formation is 98 m thick at Bow Island, 65 m at Lethbridge, Alberta, and up  
367 to 130 m in the Sweetgrass Hills area (Williams and Dyer 1930; Tuck 1993). A thin horizon of  
368 chert pebbles is present at the base of the unit. The formation was deposited during an extensive  
369 Late Cretaceous transgression episode; however the sea invasion did not reach the western part of  
370 the Sweetgrass Arch (Stebinger 1917b; Williams and Dyer 1930). Therefore, the tongue of  
371 marine shale progressively thins to zero westward. Where the Claggett/Pakowki Formations  
372 pinch out, the Milk River Formation is directly overlain by the Judith River Formation/Belly  
373 River Group. The top of the Pakowki Formation is equivalent to the top of the Lea Park  
374 Formation in central Alberta (Williams and Dyer 1930; Meijer-Drees and Mhyr 1981; Dawson et  
375 al 1994). The Pakowki/Claggett Formation effectively constitutes a regional aquitard (Fig. 6); the  
376 hydraulic conductivity of the Pakowki/Claggett Formation is in the order of  $10^{-11}$  m/s (Toth and  
377 Corbet 1986; Anna 2011).

378

### 379 **Two Medicine Formation**

380 The non-marine Two Medicine Formation of Late Cretaceous age outcrops in northwestern  
381 Montana. This unit consists of mudstones and sandstones and is about 600 m thick (Lorenz  
382 1981). West of the Sweetgrass Arch, the Two Medicine Formation includes the equivalent upper  
383 part of Eagle Formation (i.e. Deadhorse Coulee Member equivalent), the poorly recognizable  
384 Claggett Shale and the Judith River Formation (Pierce and Hunt 1937; Zimmerman 1967; Gill  
385 and Cobban 1973; Fig. 4). The Two Medicine Formation overlies the well-defined Virgelle  
386 Member. It is 152 m thick in the Cut Bank, Montana, area (Zimmerman 1967).

387

### 388 **Belly River/Judith River**

389 The Belly River Group (or equivalent Judith River Formation in Montana) outcrops in a large  
390 part of the study area (Fig. 5a). It represents the sequence of continental beds above the Pakowki  
391 Formation and below the Bearpaw Formation. The Belly River includes the Dinosaur Park  
392 Formation (upper part), the Oldman Formation (middle part) and Foremost Formation (lower  
393 part) (Eberth and Hamblin 1993; Hamblin 1997) . However, the upper part is only present in a  
394 limited part of the study area, covering of about 12 townships, near the Saskatchewan border  
395 (Hamblin 1997). Fig. 5a is adapted from the geological map from Okulitch (1996) in which the  
396 Dinosaur Park Formation is not represented.

397  
398 The dark shale, sandstone and coal seams of the Foremost Formation are overlain by massive  
399 yellow and grey sandstone of the Oldman Formation and thick sandstones and siltstones of the  
400 Dinosaur Park Formation. The Belly River Group/Judith River Formation is 320 m thick at  
401 Lethbridge, and is less than 182 m in northern Montana (Williams and Dyer 1930; Pierce and  
402 Hunt 1937). The Belly River Group/Judith River Formation constitutes an aquifer with a  
403 hydraulic conductivity ranging from  $9 \times 10^{-8}$  m/s to  $8.8 \times 10^{-7}$  m/s (Anna 2011).

404

#### 405 **Bearpaw Formation**

406 The Bearpaw Formation overlies the Belly River Group/Judith River Formation (or the Two  
407 Medicine Formation, west of the Sweetgrass Arch in Montana) and is made up of dark grey shale  
408 (Russell 1970). These marine strata were deposited during a subsequent Late Cretaceous  
409 transgression episode and are lithologically similar to the Pakowki Formation. In the western part  
410 of the study area, the Bearpaw Formation outcrops along a narrow north-south band, and around  
411 the Cypress Hills in south-eastern Alberta. The Bearpaw Formation is about 70 m thick in the  
412 north-western part of the study area and constitutes a regional aquitard (Tokarsky 1974) (Fig. 6).

413

## 414 **Available data and Methods**

415

### 416 **Data collection**

417 Geological data from five sources were gathered. In Alberta, the first source of geological data  
418 was from O'Connell (2014). Of the 2170 total borehole data, only the non-deviated wells were  
419 selected to develop the 3D geological model (2070 borehole data). These data from non-deviated  
420 wells contain the depth of the tops of the Alderson Member and the geological formations  
421 included in the Milk River Formation: Deadhorse Coulee Member, Virgelle Member and  
422 Telegraph Creek Member. The extent of the data is from 49° to 50.3° north latitude and from  
423 longitude -110° to -113°. Amongst the data provided by O'Connell (2014), there is an area of  
424 about 14 townships close to the border with no subsurface data for the Milk River Formation.  
425 This area is the structural crest of the Sweetgrass Arch, in which the Milk River Formation is too  
426 shallow to be covered by geophysical logs from hydrocarbon exploration boreholes.

427

428 The geological data of the units above the Milk River Formation were obtained from the Alberta  
429 Geological Survey/Alberta Energy Regulator. The elevations of the tops of Lea Park Formation  
430 (Milk River and Pakowki Formations equivalents) (unpublished data), Belly River Group  
431 (Glombick 2010) and the bedrock topography (Atkinson and Lyster 2010) in southern Alberta  
432 were added to the model.

433

434 In Montana, three sources of geological data were used: Feltis et al. (1981), described the  
435 elevation of the tops of each geological formation in north central Montana, from Jurassic to  
436 Quaternary. The geological formations in the Colorado Group to the Judith River Formation were

437 selected. The 190 resulting wells range from 48.2° to 49° north latitude and from longitude -  
438 109.6° to -112.5°.

439  
440 The second source of geological information consists of well logs of the area, provided by the  
441 Montana Geological Society (Montana Geological Society 2013). 103 scanned logs (mostly from  
442 the 1950's) located in Glacier, Toole, Liberty and Hill Counties were selected. They were  
443 digitized and converted into elevation format to be integrated to the data set. These logs contain  
444 the description of all the members that compose the Eagle Sandstone.

445  
446 About 10 control wells with altitude of the top of Judith River Aquifer in northern Montana were  
447 retrieved from a map produced by Noble et al. (1982). This map was scanned, georeferenced and  
448 converted from feet into meters.

449  
450 The top of the geological model is ground level, represented by a Digital Elevation Model (DEM)  
451 of the study area. The DEM in Alberta is from the National Topographic Data Base (NTDB) and  
452 the DEM of northern Montana is from USGS Earth Explorer. They are both at 1:50 000 scale (or  
453 1 arc second).

454  
455 The last source of data used in this study consisted of hydrogeological cross sections of the study  
456 area from Borneuf (1974), Tokarsky (1974) and a transboundary map of the bedrock geology  
457 from Okulitch et al. (1996).

458  
459 These various geological data required several steps of conversion and transformation in order to  
460 use them consistently in the building of the 3D geological model.

461

462 **Method: Data processing and unification of the geological data**

463

464 The geological data collected for this regional and transboundary study presented various formats  
465 and several spatial and stratigraphic references. Data processing was needed to obtain a  
466 consistent file containing the coordinates of a well and the associated elevations of each  
467 geological formation.

468

469 The main steps of data processing were: the conversion from feet to meters and from spatial  
470 reference NAD 27 to NAD 83, the transition from Township/Range system to latitude/longitude  
471 coordinates and the transition from depth to elevation data (the reference is the mean sea level).  
472 Additionally, the two available DEM files were merged to obtain a unique DEM covering the  
473 study area.

474

475 The various sets of geological data on both sides of the border are now homogeneous (same  
476 format, same spatial reference and same units). The next stage was to unify these data, since the  
477 main goal of the 3D geological model is to represent the geology of the study area in a unified  
478 way. Thus the geological data corresponding to equivalent layers in Alberta and Montana needed  
479 to be merged. In particular, specific work on the equivalent members of the Milk River  
480 Formation was needed to represent the Milk River Aquifer (i.e. Virgelle Member) in its entirety.

481

482 The collected data did not always present the same degree of information everywhere in the study  
483 area. In order to assemble the data sets for each geological unit, some working hypotheses had to  
484 be made following the 4 zones of the study area (Fig. 5a and Fig. 6).

485  
486 O’Connell (2014) described every member of the Milk River Formation in Alberta (i.e.  
487 Telegraph Creek, Virgelle and Deadhorse Coulee Members in zone 1 and the Alderson Member  
488 in zone 2). Feltis et al. (1981) only represented the “Eagle Sandstone” which is the Milk River  
489 Formation equivalent in northern Montana (zones 3 and 4). The constituting members of the  
490 Eagle Sandstone are not described individually. Considering these two sources of data, the degree  
491 of geological information on the Milk River Formation is higher in Alberta than in Montana. Yet,  
492 we want to represent the 3 members of the Milk River Formation in the entire study area.  
493 Consequently, the “Eagle Sandstone” layer from Feltis et al. (1981) was subdivided in 3 parts,  
494 using the detailed logs from the Montana Geological Society (2013). In zone 3, the constitutive  
495 members of the Two Medicine Formation were also described individually on the basis of the  
496 detailed logs from the Montana Geological Society (2013).

497  
498 Some of the wells from Feltis et al. (1981) were also part of the set of scanned logs from the  
499 Montana Geological Society (2013). The comparison of these wells clearly showed that what  
500 Feltis et al. (1981) called “Eagle Sandstone” actually included Telegraph Creek and Virgelle  
501 Members west of the Sweetgrass Arch (zone 3). East of the Sweetgrass Arch (zone 4), the  
502 comparison of the logs shows that an upper part (30 to 35 m thick) overlying the Virgelle  
503 Member was also included in the “Eagle Sandstone” layer. This upper part would equate to the  
504 “middle member” of Eagle, as defined by Payenberg et al. (2002), i.e. the Deadhorse Coulee  
505 member equivalent in Alberta.

506  
507 Based on these observations, it was assumed that east of the Sweetgrass Arch (zone 4), the top of  
508 the Eagle Sandstone described by Feltis et al. (1981) corresponds to the top of the Deadhorse



509 Coulee Member whereas, west of the Sweetgrass Arch (zone 3), it corresponds to the top of the  
510 Virgelle Member.

511  
512 West of the Sweetgrass Arch, the upper part of Eagle is included within the Two Medicine  
513 Formation and is not described. Therefore the question of how much of the lower part of the Two  
514 Medicine Formation is actually equivalent to the Deadhorse Coulee Member was raised.  
515 Reexamination of detailed boreholes from the Montana Geological Society (2013) west of the  
516 Sweetgrass Arch allowed the identification of the transgressive chert pebble lag surface, which  
517 represents the top of the Deadhorse Coulee. Too few descriptions delineated the chert pebbles  
518 marker bed, so eventually it was decided to designate the lower 75 m of the Two Medicine  
519 Formation as the Deadhorse Coulee equivalent. This value was based on the Lexicon definition  
520 of the Two Medicine Formation (WEBLEX Canada 2013), which mentions coarse sandstone  
521 beds in the lower 75 m interpreted to represent the fluvial channel sandstones of the Deadhorse  
522 Coulee. Rice and Cobban (1977) supported this value, citing Cobban (1955): “the sandstone is  
523 mainly in the basal 76 m, which is more or less equivalent to the middle and upper members of  
524 the Eagle Sandstone”. A dummy data point set representing the top of the Deadhorse Coulee  
525 west of the Sweetgrass arch was therefore added. Close to the international border, the thickness  
526 of the Deadhorse Coulee Member was reduced at 60 m to be consistent with the data of  
527 O’Connell (2014) in southern Alberta.

528  
529 The geological layers overlying the Eagle Sandstone in Montana (Claggett Shale and Judith  
530 River Formation) are only described east of the Sweetgrass Arch by Feltis et al. (1981) and the  
531 Montana Geological Society (2013). West of the Sweetgrass Arch, both of these geological layers  
532 are included in the Two Medicine Formation, which is not represented in the report of Feltis et al.

533 (1981), but is well described within the logs from the Montana Geological Society (2013). The  
534 Two Medicine Formation is equivalent to the upper part of Eagle, Claggett Shale and Judith  
535 River Formation (Zimmerman 1967). Thus, the top of the Two Medicine Formation was  
536 identified as the top of the Judith River Formation in zone 3.

537  
538 The top of the Pakowki Formation was assimilated into the top of Lea Park Formation in Alberta  
539 (Williams and Dyer 1930; Meijer-Drees and Mhyr 1981; Dawson et al 1994; Fig. 4). In northern  
540 Montana, the Claggett Shale is well described within the Montana Geological Society (2013)  
541 well logs and some of the wells of Feltis et al. (1981). However, there are no data describing the  
542 Claggett Shale west of Sweetgrass Arch since the marine invasion during which the  
543 Claggett/Pakowki was deposited did not extend far beyond the Arch. The Claggett Shale wedges  
544 out west of the Sweetgrass Arch, and where it pinches out, the Judith River(Belly River) strata  
545 rest on the Deadhorse Coulee Member equivalent.

546  
547 It is assumed that the top of the Belly River Group in southern Alberta (i.e. top of Oldman  
548 Formation or Dinosaur Park Formation, where present) is the stratigraphic equivalent to the top  
549 of the Judith River Formation in Montana (Russell 1970). The constitutive layers of the Belly  
550 River in Alberta are not represented in the model since the contact between the Oldman and  
551 Foremost Formations is difficult to identify (Williams and Dyer 1930).

552  
553 In Alberta, the top of the Colorado Group was considered to be the bottom of the Telegraph  
554 Creek Formation in zone 1, and the bottom of the Alderson Member in zone 2.

555

556 As mentioned previously, because of the lack of well logs in the southern Alberta outcrop area,  
557 the data in O’Connell (2014) was supplemented by creating a file describing the bedrock  
558 topography. The bedrock topography file was built by subtracting the thickness of surficial  
559 sediments from the topography of the study area; it was then assumed that the top of the bedrock  
560 corresponds to the top of the Deadhorse Coulee Formation where Deadhorse Coulee outcrops.  
561 The same assumption was made for the Pakowki and Belly River outcrop areas, following the  
562 geological map of Okulitch et al. (1996). Some manual editing was required to represent these  
563 outcrop areas with more details. For that purpose, the bedrock geological map (Okulitch et al.  
564 1996) was superimposed on the model as a reference.

565  
566 The regional unconformity between the Milk River Formation and the Alderson Member is also  
567 represented in the 3D model. In the north, northeast and east of the study area, the Telegraph  
568 Creek, Virgelle and Deadhorse Coulee members are progressively overlapped by the Alderson  
569 Member. An overlapping area is represented, based on the isopach maps of O’Connell (2014).  
570 The two sand bodies which make up the upper part of the Alderson Member (Upper Alderson  
571 Sands, Fig. 2) are not represented separately but they are included in the Alderson Member.

572  
573 Finally, it was decided that the Bearpaw Formation and the surficial sediments would be grouped  
574 in one layer in the geological model, between the top of the Belly River/Judith River and the  
575 ground level. With the geological data on both sides of the border harmonized and the  
576 stratigraphy unified, the 3D geological model could be built.

577

### 578 **Construction of the 3D unified geological model**

579

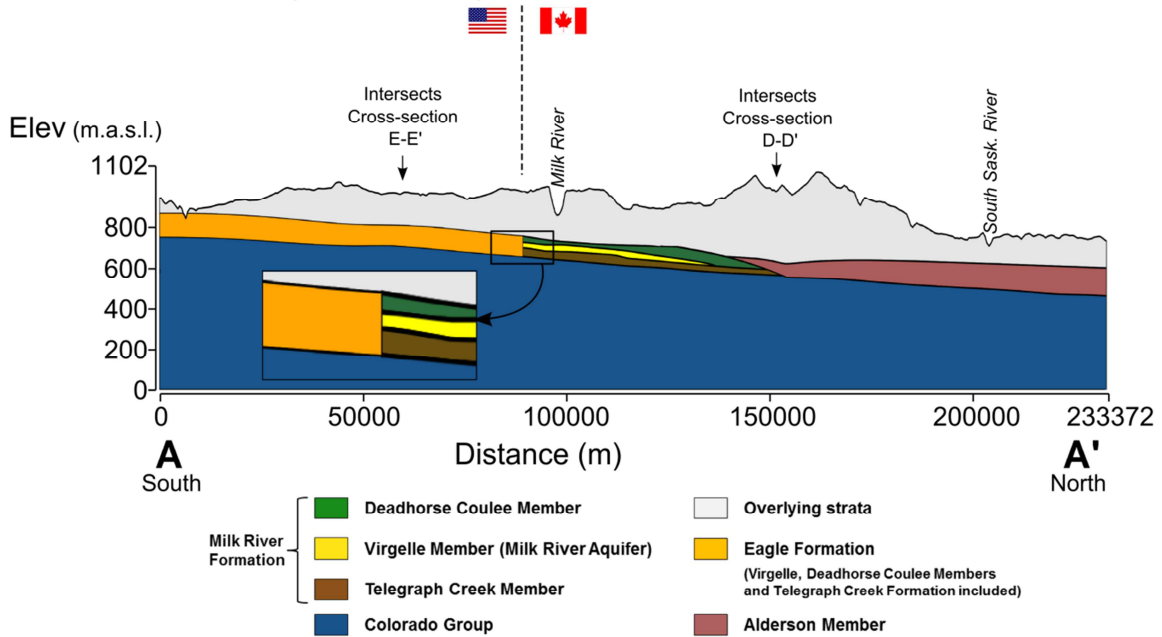
580 The software chosen to build the 3D geological model is Leapfrog Hydro<sup>®</sup> (Leapfrog Hydro  
581 2013). The approach to build the 3D geological model was to use location data (x, y, z)  
582 representing the top of the geological units. Contact surfaces were first created from these data.  
583 Then, volumes were obtained from the surfaces for which a chronology had been first  
584 determined. Leapfrog Hydro is a recent software and it was chosen mostly for its interoperability  
585 with FEFLOW (Finite Element subsurface Flow) (Diersch 2014). The 3D geological unified  
586 model will form the basis of a 3D numerical hydrogeological model of the Milk River Aquifer  
587 using FEFLOW.

588

589 An initial geological model was built as a reference and for the purpose of comparison before and  
590 after unification of the geological datasets from both sides of the international boundary. The  
591 reference model contains the data from O'Connell (2014) and Feltis et al. (1981) before the  
592 harmonization of the geological units. A south-north transboundary cross-section A-A' from this  
593 reference model is shown in Fig. 7. This section shows that the level of information was higher in  
594 southern Alberta (where Telegraph Creek, Virgelle and Deadhorse Coulee Members were  
595 represented) than in northern Montana (where the Eagle Sandstone was solely represented). The  
596 two datasets are simply placed alongside in this reference model, so the geological layers are  
597 obviously abruptly separated at the international border. The same cross-section A-A' will be  
598 shown as unified in the Results section below (Fig. 9b).

599

Cross-section A-A' (before unification)



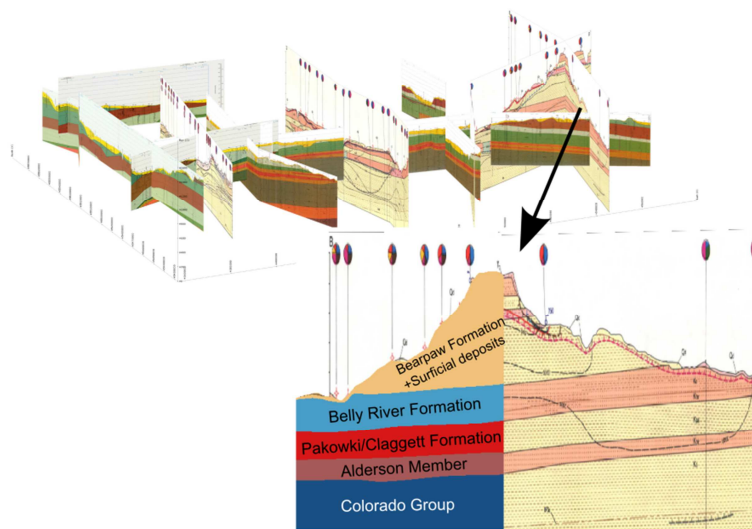
600  
 601 **Fig. 7.** Transboundary cross-section (south-north) before the unification of the Milk River/Eagle  
 602 Formation. The geological units above the Milk River/Eagle Formation are not represented in this  
 603 section. The location of the cross-section is shown in Fig. 9a.

604  
 605 In a subsequent geological model, the integration of detailed logs from Montana Geological  
 606 Society (2013) and the formulation of working hypothesis on the equivalent layers (described  
 607 above) allowed the building of the unified 3D geological model of the study area and the  
 608 separation of the Eagle Formation into three members. The datasets in the 4 zones of the study  
 609 area (Fig. 5a; Fig. 6) were finally merged to represent each geological unit.

610  
 611 Furthermore, the model was adjusted with the help of cross-sections existing for the study area.  
 612 About 15 cross-sections (Borneuf 1974; Tokarsky 1974) in Alberta were georeferenced and

613 included in the geological model. The cross sections serve as a guide and allow adjustments of  
614 the geological surfaces by manual editing within Leapfrog Hydro<sup>®</sup> (Fig. 8)

615  
616 The construction of a sloping surface was needed to represent the regional unconformity between  
617 the Milk River Formation and the Alderson Member. This surface was built by selecting points  
618 from the Alderson Member dataset, which make up a plateau overlapping Virgelle and  
619 Deadhorse Coulee Member.

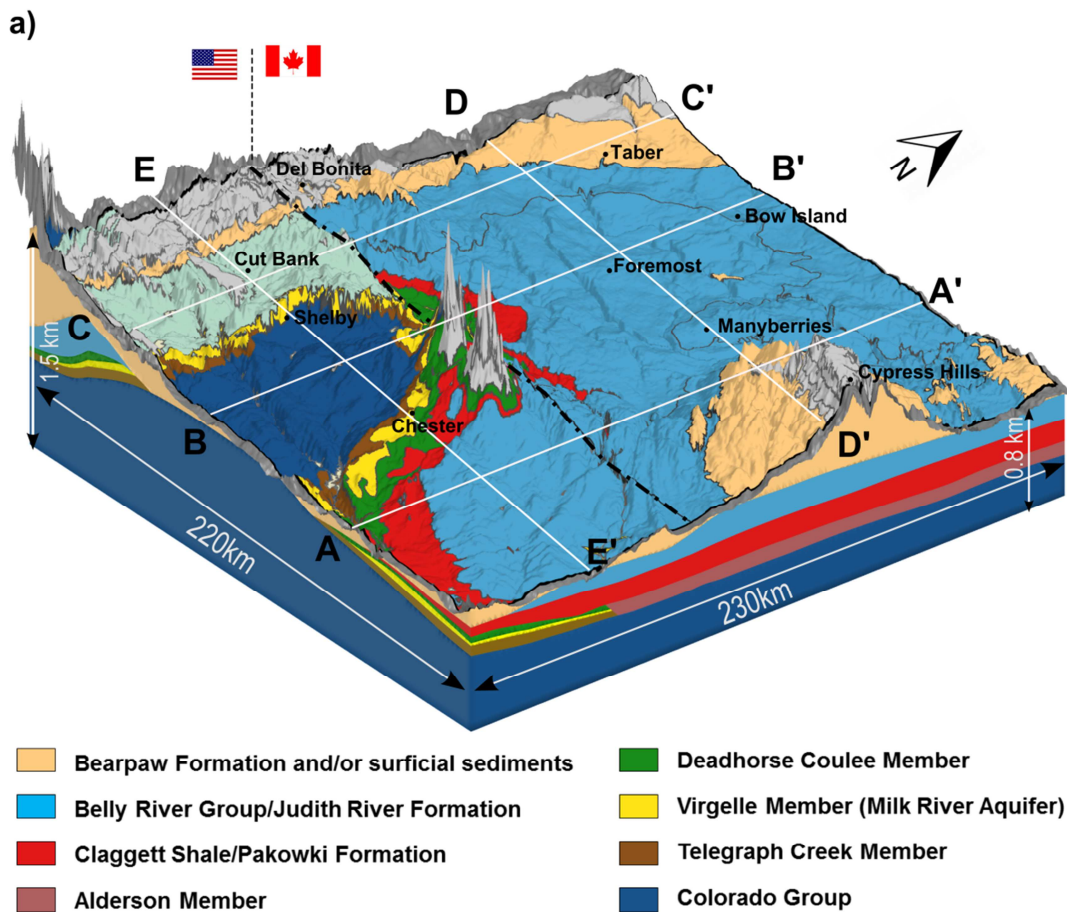


620  
621 **Fig. 8.** Examples of geo-referenced cross-sections included in the 3D geological model.

622  
623 **Results**

624  
625 The unified 3D geological model covers 50 000 km<sup>2</sup>. The geological units represented in the  
626 model are (in ascending order): Colorado Group, Telegraph Creek Member, Virgelle Member,  
627 Deadhorse Coulee Member, Alderson Member, Claggett Shale/Pakowki Formation, Belly River  
628 Group/Judith River Formation, Bearpaw Formation and surficial sediments (undivided) (Fig. 9).

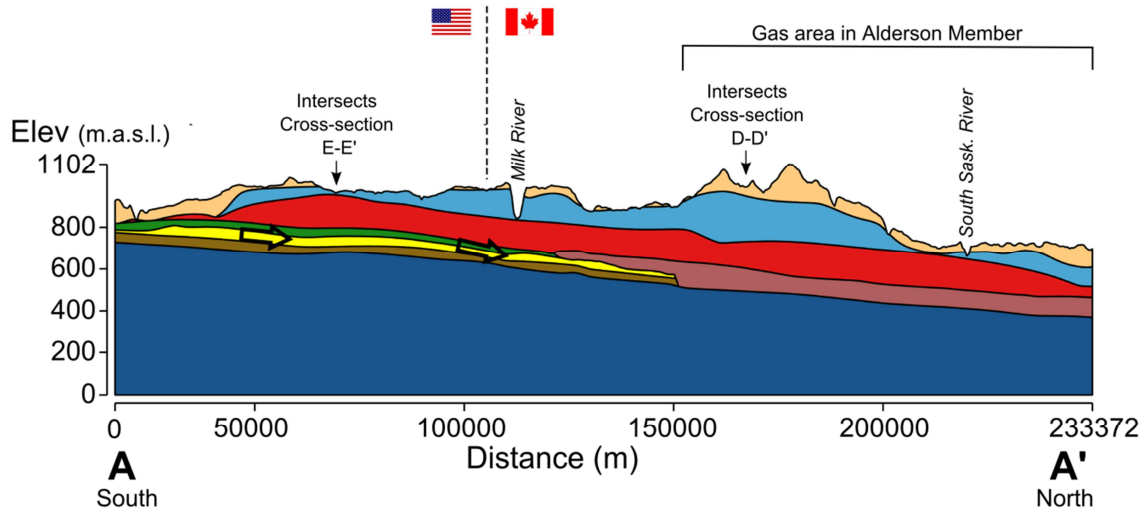
629 The locations of sections through the geological model are shown on Fig. 9a and these sections  
 630 are illustrated in Fig. 9b. The general groundwater flow direction is indicated on the cross-  
 631 sections, based on previous work (Meyboom 1960; Zimmerman 1967; Levings 1982; Tuck 1993;  
 632 AGRA Earth and Environmental Limited 1998). The area in which the Alderson Member is no  
 633 longer water-bearing but gas-bearing is indicated on the cross-sections A-A', B-B' and D-D'. The  
 634 bedrock geological map is superimposed on the model (Fig. 9a). Note that the Two Medicine  
 635 Formation is shown on the geological map; however the 3D model details all the constitutive  
 636 members of this formation, especially as shown on cross-section E-E'.



637

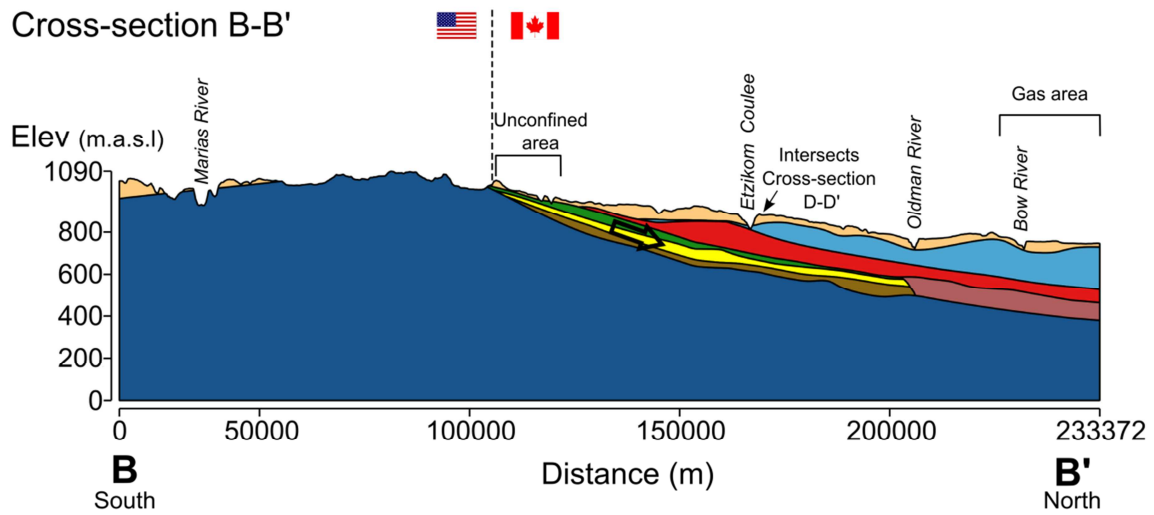
b)

Cross-section A-A' (after unification)



638

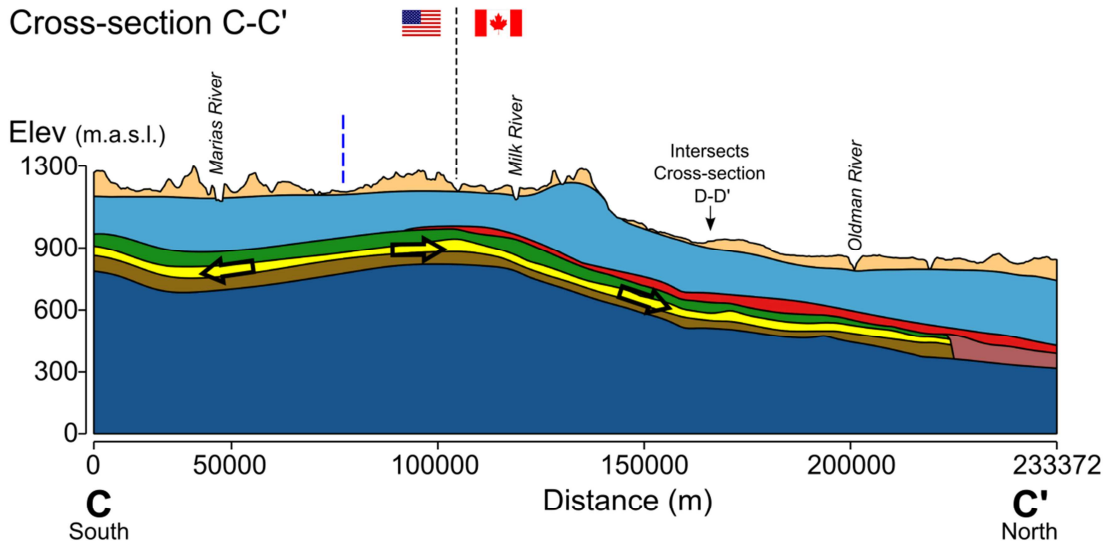
Cross-section B-B'



639

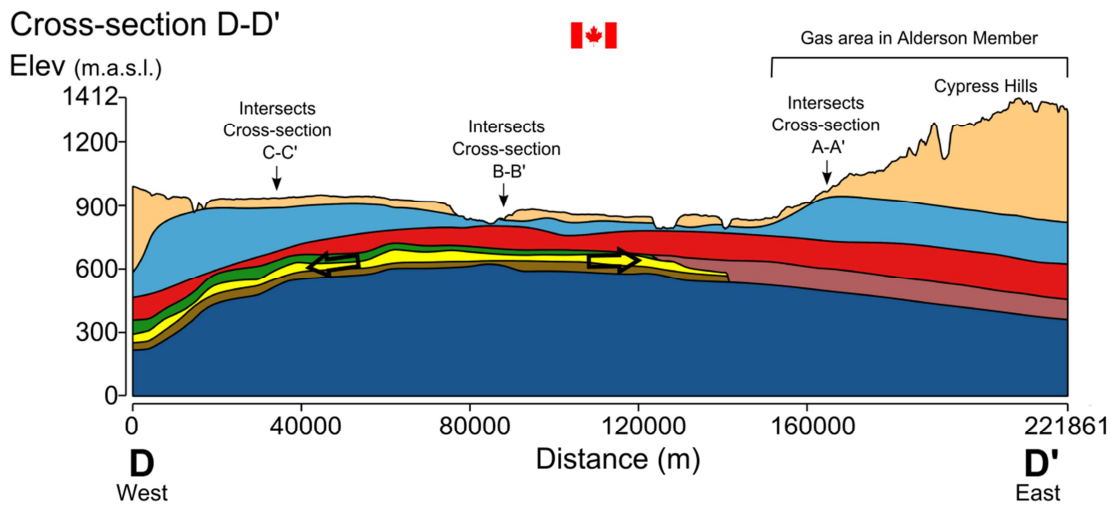


### Cross-section C-C'



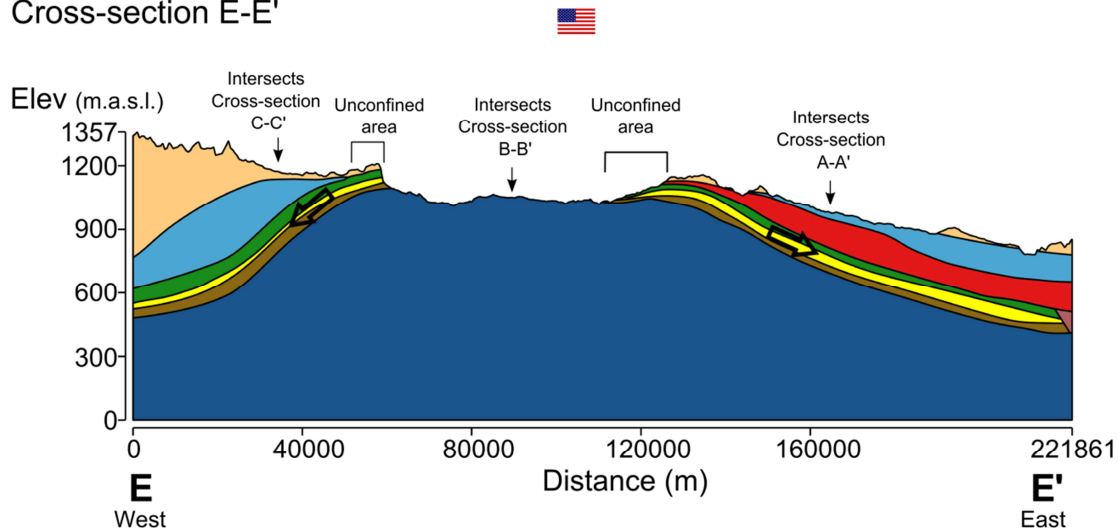
640 --- Groundwater divide after Zimmerman (1967)

### Cross-section D-D'



641

### Cross-section E-E'



642  
 643 **Fig. 9.** 3D unified geological model of the Milk River Aquifer and other geological units. **a)** 3D  
 644 unified block showing the locations of cross-sections. Vertical exaggeration factor is 50. Cross-  
 645 sections A-A', B-B' and C-C' are transboundary, cross-section D-D' is located in southern  
 646 Alberta and cross-section E-E' is located in northern Montana. **b)** Cross-sections through the 3D  
 647 geological model. The black arrows represent the general groundwater flow directions. Legend  
 648 for block diagram (Fig. 9a) and cross-sections (Fig. 9b).

649  
 650  
 651 Cross section A-A' shows that the Milk River Formation is represented on both sides of the  
 652 border with the same degree of information in the unified model (Fig. 9b). The three members  
 653 composing the Milk River Formation, especially the Milk River Aquifer (i.e. Virgelle Member),  
 654 are continuously represented from southern Alberta to north-central Montana. In Montana, the  
 655 Eagle Formation was divided in three members on the basis of detailed well logs and the working  
 656 hypothesis described previously. The Milk River Formation dips continuously from northern  
 657 Montana to southern Alberta. It does not subcrop in this section. The regional unconformity

658 between the traditional 3 members of the Milk River Formation and the Alderson Member is  
659 represented as an area of overlap. Groundwater flow is directed to the north, from northern  
660 Montana to Alberta. The overlapping area of the Alderson Member corresponds to the Upper  
661 Alderson Sands. This part is still water-bearing. However, farther north, the Alderson Member is  
662 gas-bearing.

663  
664 Cross-section B-B' (Fig. 9b) shows a steeper slope of the Milk River Formation from the  
665 international border to the north. The Colorado Group outcrops in northern Montana. A  
666 subcrop/outcrop area of the Milk River Formation at the border indicates unconfined conditions  
667 of the aquifer corresponding to a recharge area.

668  
669 West of the Sweetgrass Arch, cross-section C-C' (Fig. 9b) shows that the Claggett(Pakowki)  
670 Formation pinches out in northwestern Montana. In Alberta, the Milk River Formation is overlain  
671 by the Pakowki (Claggett) Formation and the thick Belly River Group (Judith River Formation),  
672 whereas in Montana the Judith River Formation directly overlies the Milk River Formation  
673 equivalents. The Milk River Aquifer is confined and the general groundwater flow is from south  
674 to north, except in the vicinity of Cut Bank (Fig. 5a) where it is directed to the south.

675  
676  
677 Cross-section D-D' (Fig. 9b) is located in Alberta; it shows the gentle antiformal geometry of the  
678 Milk River strata dipping eastward and westward as well as the overlap of the Alderson Member  
679 to the east. The Milk River Formation is overlain by 20 to 160 m of Pakowki(Claggett)  
680 Formation and 20 to 200 m of Belly River Group (Judith River Formation). Groundwater flows

681 to the north (perpendicular to the cross-section plan), as well as to the east and west following the  
682 aquifer elevation.

683  
684 Cross-section E-E' (Fig. 9b) is located in northern Montana in the vicinity of the Sweetgrass  
685 Arch axis. The large outcrop of the Colorado Group is represented. The Milk River Formation  
686 equivalents dip to the east and to the west on both sides of the Sweetgrass Arch. The Claggett  
687 (Pakowki) Formation is not present west of the Sweetgrass Arch, but it overlies the Eagle  
688 Formation in the east. There are two subcrop areas of the Milk River Formation equivalents  
689 which correspond to the east and west outcrop bands described above in the bedrock geological  
690 map (Fig. 5a) and the Geological Section. Therefore, the Milk River Aquifer is under unconfined  
691 conditions in these areas that represent recharge zones. Groundwater flow is directed to the south  
692 (perpendicular to the cross-section plan), as well as west and east from the subcrop areas.

693  
694 The illustration of the 3D unified geological model is complemented by several elevation and  
695 isopach maps of some of the key geological units (Fig. 10). These maps are derived from the 3D  
696 geological model which comes from the various sources previously described in the "Data  
697 Collection" section. The elevation of the top of the Milk River Formation (including Telegraph  
698 Creek, Virgelle, Deadhorse Coulee and Alderson Members) is shown in Fig. 10a. The Milk River  
699 Formation dips to the north, east and west from the outcrop area near the international border  
700 following a semi-radial pattern. This pattern is continuous through the international border as the  
701 Milk River formation dips to the east, south-east and north from the Sweetgrass Hills and to the  
702 west, in the area west of the Sweetgrass Arch. The top of the Milk River Formation reaches about  
703 1100 m at the border and decreases to 500 m in the northern part of the study area, about 130 km  
704 north of the border. Therefore, the average slope is around 4.7 m/km. In Montana, the dip of the

705 Milk River Formation equivalents is more pronounced; the top of the formation attains 1100 m  
706 around the Sweetgrass Hills and drops to 550 m in the eastern part of the study area, the average  
707 slope is therefore 9 m/km. West of the Sweetgrass Arch, the top of the Milk River Formation  
708 equivalents drops from 1150 m to 600 m in less than 50 km, so the mean slope is around 13  
709 m/km. The thickness of the Milk River Formation ranges from 0 to 140 m in Alberta and from 0  
710 to 200 m in Montana. In the subcrop areas, the formation is 0 to 60 m thick.

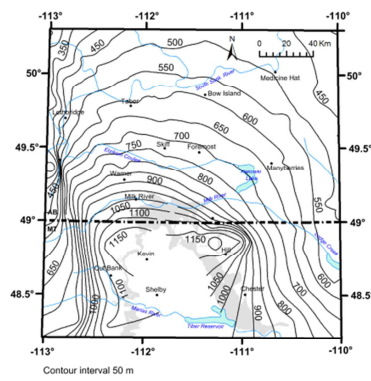
711  
712 The elevation of the Virgelle Member top is shown in Fig. 10b. As the middle member of the  
713 Milk River Formation, it follows the same dipping configuration. The unit thickness ranges  
714 between 0 to 60 m in the study area. The Virgelle Member thins towards the deposition limit,  
715 which extends from the northwest corner of the study area to the east, near the international  
716 border. The extent of the Alderson and Virgelle Members (Fig. 10c) are mutually exclusive  
717 except for an overlapping area in the vicinity of the deposition limit of Virgelle where the  
718 Alderson Member is 0 to 15 m thick. The Alderson Member thickens to the north and east and  
719 reaches a thickness of 95 m in the northeast corner of the study area.

720  
721 The Milk River Formation is confined by the overlying Pakowki/Claggett Formation which  
722 follows the general semi-radial dip of the Milk River Formation (Fig. 10d). The elevation of the  
723 top of the Pakowki/Claggett Formation is about 900 m north of the subcrop area of the Milk  
724 River Formation. It reaches an elevation of 550 m approximately 120 km farther north. The  
725 thickness of this unit ranges from 0 to 160 m; the unit thins from east to west.

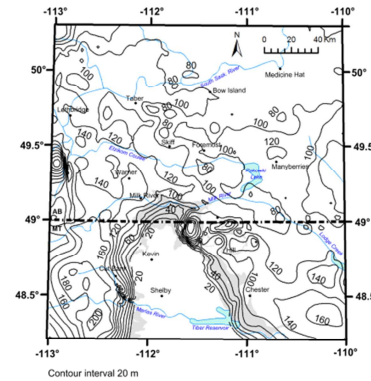
726 The Claggett Shale is not present west of the Sweetgrass arch; it might not be present in the  
727 south-west area close to the border in Alberta. Due to the lack of data in the area where Claggett  
728 wedges out, the contours are approximate.

**a) Milk River Formation and Montana equivalents**

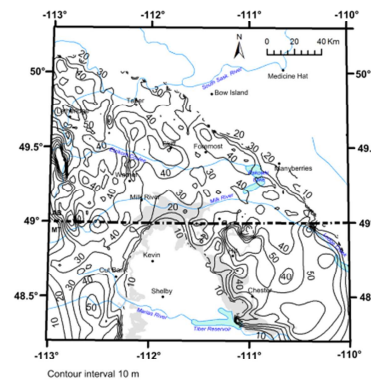
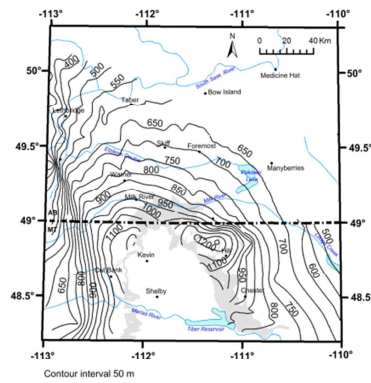
Elevation (z)



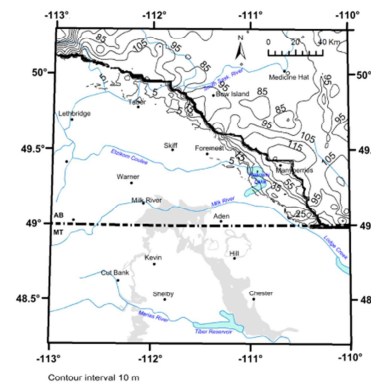
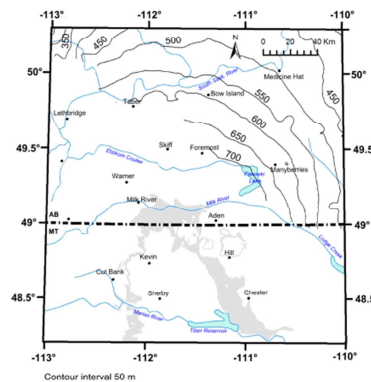
Thickness (e)



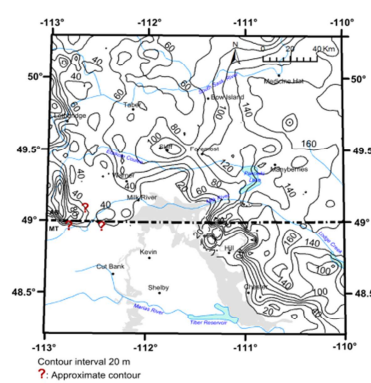
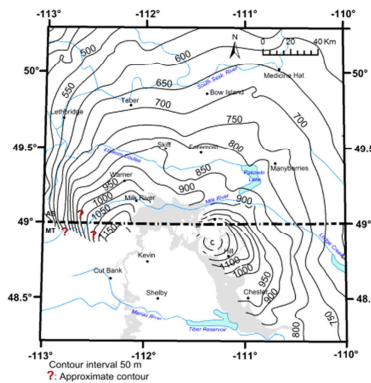
**b) Virgelle Member**



**c) Alderson Member**



**d) Pakowki/Claggett Formation**



Subcrop area of Milk River Formation

730 **Fig. 10.** Unified elevation contours of the top of several geological units (left) and isopach maps  
731 (right). The elevation reference is the mean sea level and the thickness is in meters. These maps  
732 are derived from the 3D geological model which comes from the various sources previously  
733 described in the “Data Collection” section. **a)** Elevation and thickness of the Milk River  
734 Formation (including Telegraph Creek, Virgelle, Deadhorse Coulee and Alderson Members). **b)**  
735 Elevation and thickness of the Virgelle Member. **c)** Elevation and thickness of the Alderson  
736 Member. **d)** Elevation and thickness of the Pakowki/Claggett Formation.

737

738

## 739 **Discussion**

740

### 741 **Changes and confirmations in the geological context**

742 The cross-sections (Fig. 9) and the elevation and isopach maps (Fig. 10) corroborate previous  
743 studies of the Milk River Aquifer. In particular, the semi-radial pattern of the elevation of the  
744 Milk River Formation top shown in Fig. 10a is in agreement with Meyboom (1960) and others in  
745 southern Alberta. Besides, the unified 3D geological model completes this observation by  
746 showing that the semi-radial pattern continues into northern Montana.

747

748 The 3D geological model represents the subcrop areas of the Milk River Formation near the  
749 Sweetgrass Hills and near the international border, which are also traditionally well-documented.  
750 However, the model also indicates two additional subcrop areas on both sides of the Sweetgrass  
751 Arch in Montana.

752

753 This study represents the Alderson Member distinctly from the other three members of the Milk  
754 River Formation, separated by a regional unconformity. Previously, the interpreted gradual facies  
755 change of the Milk River Formation was only conceptually represented but it was not included in  
756 a 3D geological model.

757  
758 Furthermore, the 3D unified geological model includes the area west of the Sweetgrass Arch in  
759 Montana (zone 3). This area was little documented in previous Milk River Aquifer studies. Yet,  
760 this is an important area since it hosts the southwestern part of the Milk River Aquifer. Thus,  
761 consideration of this zone allows a complete geological representation of the aquifer. The  
762 originality of the present work also lies in the representation of the three members of the Milk  
763 River Formation continuously across the international border (especially the middle Virgelle  
764 Member which is the Milk River Aquifer), as well as the overlying and underlying strata. The 3D  
765 geological model thus represents the natural limits of the Milk River Aquifer, not influenced by  
766 jurisdictional boundaries.

767  
768 **Hydrogeological implications of the geological model**

769 The elevation contour maps and cross-sections provide a better appreciation of the confined  
770 conditions of the aquifer. The Milk River Formation dips to the north from the outcrop areas in  
771 the Sweetgrass Hills and near the border in southern Alberta (zones 1 and 2). It dips to the south  
772 and east in zone 3 and to the south and west in zone 4. The Milk River Aquifer is therefore  
773 confined throughout most of the study area by the overlying Pakowki/Claggett Formations. The  
774 recharge areas of the Milk River aquifer have been traditionally identified as the outcrop areas  
775 near the border in southern Alberta and around the Sweetgrass Hills in Montana. However, the



776 Milk River Formation and most importantly the Virgelle Member, subcrop also along both sides  
777 of the Sweetgrass Arch in Montana (Fig. 4a; Fig. 9a), which could also represent recharge areas.

778  
779 The Milk River Aquifer is limited to the north, north-east, and east by the erosional unconformity  
780 surface which is overlapped by its lateral equivalent the Alderson Member. Since Alderson  
781 Member is included in the Milk River Formation (O'Connell 2014; Payenberg et al. 2003) and its  
782 upper part (Upper Alderson Sand) is water-bearing, it has been included in the delineation of the  
783 Milk River Aquifer in Southern Alberta by Printz (2004). However, Printz (2004) did not  
784 mention the presence of the Alderson Member, unconformably incised into the older members of  
785 the Milk River Formation. O'Connell (2014) considered that the Milk River Aquifer consists of  
786 two regional sand units within the Milk River Formation: the Virgelle Sand and the Upper  
787 Alderson Sands. However, Alderson Member is much younger than the three other members of  
788 the Milk River Formation (Payenberg et al. 2002). In addition, unlike the Virgelle Member which  
789 is transboundary and continuous through the international border, the Alderson member is not  
790 present in Montana where its chronostratigraphic equivalent is located south of the study area  
791 (Payenberg et al. 2003). These differences lead us to consider the Alderson Member as distinct  
792 from the other members of the Milk River Formation. Thus, the present study limits the Milk  
793 River transboundary aquifer to the transboundary Virgelle Member only.

794  
795 The Milk River Aquifer is limited by the Medicine Hat (Milk River) gas field hosted by the  
796 Alderson Member, north, north-east and east of the study area in Alberta. The Tiger Ridge gas  
797 field located near the city of Havre (near the Bears Paw Mountains in Montana) represents the  
798 south-eastern boundary of the aquifer. In Alberta, the Alderson Member represented in the 3D  
799 model, overlaps the lower members of the Milk River Formation. There is no equivalent

800 relationship for the Eagle Formation which hosts the Bearpaw gas field in Montana. For the  
801 purpose of the 3D geological model, the Marias River is used as the southern limit of the Aquifer  
802 in Montana. Although the Milk River/Eagle Formation extends farther south in Montana, the  
803 Marias River physiographic limit has been chosen with respect to the future hydrogeological  
804 model and considering the available data. The western limit of the aquifer corresponds to the  
805 westernmost area in which water wells have been completed in the Virgelle Member. However,  
806 the 3D geological model shows that Virgelle Member exists farther west around the longitude -  
807 113° (Fig. 2). No water well is completed in this unit beyond the proposed limit because the  
808 aquifer is too deep (Stantec 2002). Thus, the western limit of the aquifer was based on the extent  
809 of Virgelle Member water wells.

810 This work provides a comprehensive geological framework supporting the development of a  
811 conceptual model of the Milk River Aquifer. However, other components of the conceptual  
812 model still need to be assessed (groundwater-surface water interaction, discharge areas,  
813 quantification of transboundary fluxes, geochemistry, etc). Further work will address these  
814 components through the development of a unified hydrogeological numerical model of the  
815 aquifer.

816

### 817 **Model limitations**

818 The amount of available data in Alberta is much higher than in Montana. In order to improve the  
819 quality of the 3D geological model, collecting more well log descriptions within northern  
820 Montana would be necessary. Regarding the overall resolution of the regional model, the density  
821 of geological data is less than 1 datum per km<sup>2</sup>. According to the classification of geological  
822 models of Culshaw (2005), the geological model would thus be of the “Overview” type.

823

824 Simplifications were made concerning the Pakowki/Claggett volume in zone 4. Previous studies  
825 described that the transgressive strata did not reach the 112<sup>th</sup> meridian or wedges out in the  
826 vicinity of Cut Bank (Stebinger 1917a; Russell 1970). The volume was not represented west of  
827 the Sweetgrass Arch. A rough slimming of the volume was manually edited from the outcrop  
828 area of Pakowki in southern Alberta in zone 1 towards the south in zone 3, just north of the  
829 Virgelle outcrop. More geological data are needed in the northern part of zone 3 in order to  
830 represent the Pakowki equivalent and its disappearance in a more realistic way. Another  
831 simplification was made in considering that the interval between the top of the Belly River  
832 Group/Judith River Formation and land surface included the Bearpaw Formation (when present)  
833 and surficial sediments.

834  
835 The effort for unification of the Milk River Formation across the international border implied  
836 identifying its 3 constitutive members in zones 3 and 4 in Montana, where these members had not  
837 previously been defined. This could be done more easily in zone 4 than in zone 3 thanks to the  
838 correlation work of Payenberg et al. (2002). The Eagle Sandstone described by Feltis et al. (1981)  
839 included the basal Telegraph Creek, Virgelle Member and an upper part, just below the Claggett  
840 Shale, which was equated to the Deadhorse Coulee (or middle member of Eagle, following the  
841 nomenclature of Payenberg et al. 2002). The upper member of Eagle described in Choteau and  
842 Fergus Counties, Montana, does not exist in the study area, according to its depositional area map  
843 (Payenberg et al. 2003).

844  
845 In zone 3, the absence of the Claggett/Pakowki Formation makes the unification work more  
846 complex. The Telegraph Creek and Virgelle Members were positively described in well logs.  
847 However, the overlying layers were included in the Two Medicine Formation. Given the lack of

848 data in that zone, it was decided to consider the lower 75 m of the Two Medicine Formation as  
849 the Deadhorse Coulee equivalent in zone 3, following the description of the Lexicon (WEBLEX  
850 Canada 2013). However, isopach maps from O'Connell (2014) indicated that the thickness of the  
851 Deadhorse Coulee was about 60 m at the border. The thickness of the dummy Deadhorse Coulee  
852 in zone 3 was thus reduced to 60 m close to the border (about 10 km north of the border) to be  
853 consistent and continuous with O'Connell's data (2014). The default thickness of 75 m was  
854 applied in the remaining area of the geological model. This default value is approximate and  
855 might not reflect reality in all of zone 3. However this is considered as an acceptable hypothesis  
856 in this regional geological framework study.

857  
858 The integration of georeferenced cross-sections in the geological model was valuable. It allows  
859 verification of how representative the collected geological data were, and the manual editing of  
860 geological surfaces based on georeferenced cross-sections compensates for the lack of data  
861 encountered in some places (Fig. 8). This technique was particularly useful for the  
862 Claggett/Pakowki and Belly River/Judith River surfaces. Cross-sections also show that the  
863 hypothesis of equating the tops of Belly River, Pakowki or Deadhorse Coulee to the bedrock  
864 topography, where these formations outcropped, was valid. However, there is no cross-section  
865 included in the model on the USA side. Cross-sections from the geological maps of the  
866 Sweetgrass Hills and Cutbank 30'x60' quadrangles (Lopez 2001; Berg 2002) did not show a  
867 sufficient vertical exaggeration. Thus, the quality of these cross-sections in northern Montana did  
868 not allow their use for the development of the 3D geological model.

869  
870 The Upper Alderson Sands which form small lobate sand bodies was not represented in the  
871 model, due to the limitation of the modeling software. However, this unit will be represented in

872 the future hydrogeological FEFLOW model of the aquifer, with distinct hydrogeological  
873 properties from the Alderson Member.

874  
875 Leapfrog Hydro<sup>®</sup> was chosen mostly for its interoperability with FEFLOW. When using data of  
876 various formats in Leapfrog Hydro<sup>®</sup>, the geological model results in several geological models,  
877 each with a specific data format, that are assembled to make up the whole model. Yet, when  
878 exporting the model from Leapfrog Hydro<sup>®</sup> to FEFLOW, only one geological model can be  
879 selected for transfer into FEFLOW. This implies the use of data of the same format (locations file  
880 with X, Y, Z coordinates).

881  
882 The area where the Alderson Member overlaps the lower members of the Milk River Formation  
883 was difficult to represent with Leapfrog Hydro because many operations and manual editing were  
884 necessary. Other software (such as Gocad) could have provided more flexibility for that  
885 particular task. The geological model is dynamic; new data or editing of surfaces could be added  
886 in an iterative model improvement process.

887  
888 **Conclusion and Perspectives**

889  
890 Geological data on both sides of the Canada/USA border were gathered and processed in order to  
891 build a unified three-dimensional geological model of the Milk River Aquifer.

892 The main contributions of this paper are:  
893 -Our work allowed the unification of the stratigraphic framework of the Milk River  
894 Transboundary Aquifer on both sides of the Canada/USA border.

895 -There is a more rigorous delineation of the extent of the Milk River Transboundary Aquifer, and  
896 a better description of its characteristics (thickness, geometry) following its natural boundaries  
897 (not the jurisdictional boundaries).

898 -The 3D geological model provides a representation of the three members of the Milk River  
899 Formation (especially the Virgelle Member (Milk River Aquifer) and encasing units continuously  
900 through the border.

901 -The 3D geological model also provides a description of the geological boundaries of the Milk  
902 River Aquifer, imposed by the regional structure, the regional unconformity in southern Alberta,  
903 and the sealing effects of the adjacent gas fields.

904 -The 3D geological model also includes the representation of the gas-bearing Alderson Member  
905 and the regional unconformity surface which separates it from the lower members of the Milk  
906 River Formation.

907 -The implications of the geological model on the hydrogeological conditions of the aquifer have  
908 been highlighted. The hydrostratigraphic role (aquifer or aquitard) of each geological unit has  
909 been shown.

910  
911 This 3D unified geological model is a major component of a unified conceptual hydrogeological  
912 model of the Milk River Aquifer. It will form the basis for the future development of a numerical  
913 model of the Milk River Aquifer. The next stage in this study will be to propose a unified  
914 hydrogeological conceptual model of the Milk River aquifer, including boundary conditions,  
915 groundwater flow systems, and groundwater quality. The final stage will be to transfer the  
916 geological model into FEFLOW that will be used to develop a numerical groundwater flow  
917 model. This transfer of the geological model will allow hydrogeological properties to be assigned  
918 to each geological layer.

919

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933

934 **References**

935

936 AGRA Earth and Environmental Limited 1998. Evaluation of depletion of the Milk River  
937 aquifer. AGRA Earth & Environmental, Edmonton.

938

939 Alberta Innovates Technology Future 2010. Milk River Aquifer Hydrogeology Report. Internal  
940 Report for the Geological Survey of Canada, Quebec Division.

941

942 Anna, L.O. 2011. Effects of groundwater flow on the distribution of biogenic gas in parts of the  
943 northern Great Plains of Canada and United States. USGS Scientific Investigations Report:  
944 2010-5251, 24p.  
945

946 Atkinson, N., and Lyster, S. 2010. Bedrock Topography of Alberta, Canada. Energy Resources  
947 Conservation Board, ERCB/AGS Map, 550.  
948

949 Berg, R.B. 2002. Geologic map of the Cut Bank 30' x 60' quadrangle. Montana Bureau of Mines  
950 and Geology: Open-File Report 454, 10 p., 1 sheet.  
951

952 Borneuf, D.M. 1974. Hydrogeological map Foremost, Alberta: NTS mapsheets 72E, Alberta  
953 Research Council, Edmonton, Alberta.  
954

955 Borneuf, D.M. 1976. Hydrogeology of the Foremost Area. Alberta, Alberta Research Council,  
956 Edmonton, Alberta.  
957

958 Cobban, W.A. 1950. Telegraph Creek Formation of Sweetgrass Arch, North-Central Montana:  
959 GEOLOGICAL NOTES. AAPG Bulletin **34**:1899–1900.  
960

961 Cobban, W.A. 1955. Cretaceous rocks of northwestern Montana, in Billings Geol. Soc.  
962 Guidebook 6th Ann. Field Conf., Sweetgrass Arch-Disturbed Belt, Montana, p. 107-119.  
963



964 Cobban, W.A., Erdmann, C.E., Lemke, R.W., and Maughan, E.K. 1976. Type sections and  
965 stratigraphy of the Blackleaf and Marias River Formations (Cretaceous) of the Sweetgrass Arch,  
966 Montana. USGS Professional Paper 974: 36–60.  
967

968 Culshaw, M.G. 2005. From concept towards reality: developing the attributed 3D geological  
969 model of the shallow subsurface. *Quarterly Journal of Engineering Geology and*  
970 *Hydrogeology* **38**:231–284. doi: 10.1144/1470-9236/04-072.  
971

972 Diersch, H.-J.G. 2014. FEFLOW: finite element modeling of flow, mass and heat transport in  
973 porous and fractured media. Springer, 996 p.  
974

975 Dawson, F.M., Evans, C.G., Marsh, R., and Richardson, R. 1994. Uppermost cretaceous and  
976 tertiary strata of the western Canada sedimentary basin; in *Geological Atlas of the Western*  
977 *Canada Sedimentary Basin*, G.D. Mossop and I. Shetsen (comp.), Canadian Society of Petroleum  
978 Geologists and Alberta Research Council.  
979

980 Dowling, D.B. 1915. Water supply, southeastern Alberta. Summary Report 1915; Geological  
981 Survey of Canada, Summary Report (1915), 1916; p. 102-110.  
982

983 Dowling, D.B. 1917. Southern plains of Alberta. Geological Survey of Canada, “A” Series Map  
984 187A, 1917; 1 sheet, doi:10.4095/106740.  
985

986 Dyer, W.S. and Williams, M.Y. 1930. Geology of southern Alberta and southwestern  
987 Saskatchewan. Memoir 163, Geological Survey of Canada.

988

989 Eberth, D.A., and Hamblin, A.P. 1993. Tectonic, stratigraphic, and sedimentologic significance  
990 of a regional discontinuity in the upper Judith River Group (Belly River wedge) of southern  
991 Alberta, Saskatchewan, and northern Montana. *Can J Earth Sci* **30**:174–200. doi: 10.1139/e93-  
992 016.

993

994 Evans, C. S. 1931. Milk River area and the Red Coulee Oil Field, Alberta. Geological Survey of  
995 Canada, Summary Report (1930), p. 1–30.

996

997 Feltis, R.D., Lewis, B.D., Frasure, R.L., Rioux, R.P., Jauhola, C.A., and Hotchkiss, W.R. 1981.  
998 Selected geologic data from the Northern Great Plains area of Montana. USGS Open-File  
999 Report 81-415.

1000

1001 Gautier, D., and Rice, D. 1982. Conventional and Low-Permeability Reservoirs of Shallow Gas  
1002 in the Northern Great Plains. *Journal of Petroleum Technology*. doi: 10.2118/9846-PA.

1003

1004 Gill, J.R., and Cobban, W.A. 1973. Stratigraphy and geologic history of the Montana group and  
1005 equivalent rocks, Montana, Wyoming, and North and South Dakota. USGS Professional Paper  
1006 776, 25p.

1007

1008 Glombick, P.M. 2010. Top of the Belly River Group in the Alberta Plains: Subsurface  
1009 Stratigraphic Picks and Modelled Surface. Open file Report 2010-10. Energy Resources  
1010 Conservation Board, ERCB/AGS.

1011

1012 Government of Alberta, Alberta water for life 2006. Water conservation and allocation guideline  
1013 for oil field injection.

1014

1015 Hamblin, A.P. 1997. Regional Distribution and Dispersal of the Dinosaur Park Formation, Belly  
1016 River Group, Surface and Subsurface of Southern Alberta. *Bulletin of Canadian Petroleum  
1017 Geology* **45**:377–399.

1018

1019 Hamblin, A.P., and Lee P.J. 1997. Uppermost Cretaceous, post-Colorado Group gas resources of  
1020 the Western Canada Sedimentary Basin, Interior Plains. Geological Survey of Canada. *Bulletin  
1021* 518.

1022

1023 Hendry, M.J., and Schwartz, F.W. 1988. An alternative view on the origin of chemical and  
1024 isotopic patterns in groundwater from the Milk River Aquifer, Canada. *Water Resour Res*  
1025 **24**:1747–1763. doi: 10.1029/WR024i010p01747.

1026

1027 Hendry J., Schwartz, F.W., and Robertson, C. 1991. Hydrogeology and hydrochemistry of the  
1028 Milk River aquifer system, Alberta, Canada: a review. *Applied Geochemistry* **6**:369–380. doi:  
1029 10.1016/0883-2927(91)90037-P.

1030

1031 LEAPFROG Hydro<sup>®</sup> (2013) ARANZ Geo, <http://www.leapfrog3d.com>.

1032

1033 Levings, G.W. 1982. Potentiometric-surface map of water in the Eagle Sandstone and equivalent  
1034 units in the Northern Great Plains area of Montana. USGS Open-File Report 82-565.

1035

1036 Lopez, D.A. 2001. Geologic map of the Sweet Grass Hills 30' x 60' quadrangle, north-central  
1037 Montana, Montana Bureau of Mines and Geology Open-File Report 443, 4 p., 1 sheet, 1:100 000.  
1038  
1039 Lorenz, J.C. 1981. Sedimentary and Tectonic History of the Two Medicine Formation, Late  
1040 Cretaceous (Campanian), Northwestern Montana. Princeton University.  
1041  
1042 Meijer-Drees, N.C., and Mhyr, D.W. 1981. The Upper Cretaceous Milk River and Lea Park  
1043 Formations in Southeastern Alberta. *Bulletin of Canadian Petroleum Geology* **29**:42–74.  
1044  
1045 Meyboom, P. 1960. Geology and groundwater resources of the Milk River sandstone in southern  
1046 Alberta. Research Council of Alberta.  
1047  
1048 Meyer, R. 1998. Sedimentology, petrology and permeability characterization of the Upper  
1049 Cretaceous Virgelle Member, Milk River Formation, Writing-on-Stone Provincial Park, Alberta,  
1050 Canada. University of Calgary.  
1051  
1052 Meyer, R., and Krause, F.F. 2006. Permeability anisotropy and heterogeneity of a sandstone  
1053 reservoir analogue: An estuarine to shoreface depositional system in the Virgelle Member, Milk  
1054 River Formation, Writing-on-Stone Provincial Park, southern Alberta. *Bulletin of Canadian  
1055 Petroleum Geology* **54**:301–318.  
1056  
1057 Montana Geological Society 2013. Northwest Geologic Service Sample Logs. Northern Rockies  
1058 Geological Data Center.  
1059

1060 Noble, R., Bergantino, R., Patton, T.W., Sholes, B.C., Daniel, F. and Scofield, J. 1982. Altitude  
1061 in Feet on Top of the Judith River Aquifer, *in* Occurrence and characteristics of ground water  
1062 in Montana. Volume 1. The Great Plains Region.  
1063

1064 O'Connell, S., 2014. The Milk River Transboundary Aquifer in Southern Alberta; Geological  
1065 Survey of Canada, Open File 7751, 1 .zip file. doi:10.4095/295603.  
1066

1067 Okulitch, A.V., Lopez, D.A., and Jerzykiewicz, T. 1996. Bedrock geology, Lethbridge, Alberta-  
1068 Saskatchewan-Montana, Geological Survey of Canada, National Earth Science Series,  
1069 Geological Atlas no. NM-12-G, doi:10.4095/208987.  
1070

1071 Payenberg, T.H.D. 2002a. Integration of the Alderson Member in southwestern Saskatchewan  
1072 into a litho- and chronostratigraphic framework for the Milk River/Eagle coastline in southern  
1073 Alberta and north-central Montana. *In*: Summary of Investigations 2002, Volume 1,  
1074 Saskatchewan Geological Survey, Sask. Energy Mines, Misc. Rep. 2002-4.1.  
1075

1076 Payenberg, T.H.D. 2002b. Litho-, Chrono- and Allostratigraphy of the Santonian to Campanian  
1077 Milk River and Eagle Formations in Southern Alberta and North-central Montana :  
1078 Implications for Differential Subsidence in the Western Interior Foreland Basin. Thesis  
1079 (Ph.D.) University of Toronto.  
1080

1081 Payenberg, T.H.D., Miall, A.D., and Braman, D.R. 2001. Sequence stratigraphy of the Milk  
1082 River Formation in southern Alberta and Eagle Formation in northern Montana. Rock the  
1083 foundation Convention. Canadian Society of Petroleum Geologists.

1084

1085 Payenberg, T.H.D., Braman, D.R., Davis, D.W., and Miall, A.D. 2002. Litho-and  
1086 chronostratigraphic relationships of the Santonian-Campanian Milk River Formation in  
1087 southern Alberta and Eagle Formation in Montana utilising stratigraphy, U-Pb  
1088 geochronology, and palynology. *Canadian Journal of Earth Sciences* **39**:1553–1577.

1089

1090 Payenberg, T.H.D., Braman, D.R., and Miall, A.D. 2003. Depositional environments and  
1091 stratigraphic architecture of the Late Cretaceous Milk River and Eagle formations, southern  
1092 Alberta and north-central Montana: relationships to shallow biogenic gas. *Bulletin of*  
1093 *Canadian Petroleum Geology* **51**:155–176.

1094

1095 Pierce, W.G., and Hunt, C.B. 1937. Contributions to economic geology (short papers and  
1096 preliminary reports), 1934-36. Geology and mineral resources of north-central Chouteau,  
1097 western Hill, and eastern Liberty counties, Montana. United States Geological Survey.

1098

1099 Printz, J. 2004. Milk River Aquifer Reclamation & Conservation Program 1999-2004 Summary  
1100 Report.

1101

1102 Rice, D.D. 1976. Depositional environments of the Eagle Sandstone, north-central Montana-- an  
1103 aid for hydrocarbon exploration (modified from a talk presented to the Rocky Mountain Section  
1104 AAPG-SEPM Meeting, Billings, Montana, March 30, 1976. United States Geological Survey.

1105

1106 Rice, D.D. 1980. Coastal and deltaic sedimentation of Upper Cretaceous Eagle Sandstone;  
1107 relation to shallow gas accumulations, North-central Montana. *AAPG Bulletin* **64**:316–338.

1108

1109 Rice, D.D., and Cobban, W.A. 1977. Cretaceous stratigraphy of the Glacier National Park area,  
1110 northwestern Montana. *Bulletin of Canadian Petroleum Geology* **25**:828–841.

1111

1112 Robertson, C. 1988. Potential impact of subsurface irrigation return flow on a portion of the Milk  
1113 River and Milk River Aquifer in southern Alberta . University of Alberta Dept of Geology.

1114

1115 Ross, M., Parent, M., and Lefebvre, R. 2005. 3D geologic framework models for regional  
1116 hydrogeology and land-use management: a case study from a Quaternary basin of  
1117 southwestern Quebec, Canada. *Hydrogeology Journal* **13**:690–707. doi: 10.1007/s10040-004-  
1118 0365-x.

1119

1120 Russell, L.S., and Landes, R.W. 1940. Geology of the southern Alberta Plains, Memoir 221,  
1121 Geological Survey of Canada, 219 p.

1122

1123 Russell, L.S. 1970. Correlation of the Upper Cretaceous Montana Group between southern  
1124 Alberta and Montana. *Canadian Journal of Earth Sciences* **7**:1099.

1125

1126 Schwartz, F.W., and Muehlenbachs, K. 1979. Isotope and ion geochemistry of groundwaters in  
1127 the Milk River Aquifer, Alberta. *Water Resources Research* **15**:259–268. doi:  
1128 10.1029/WR015i002p00259.

1129

1130 Stantec 2002. Regional groundwater assessment of potable groundwater in county of Warner  
1131 No5, Alberta. County of Warner No5.

1132

1133 Stanton, T.W., Hatcher, J.B., and Knowlton, F.H. 1905. Geology and paleontology of the Judith  
1134 River beds. United States Geological Survey.

1135

1136 Stebinger, E. 1915. The Montana Group of northwestern Montana. Shorter contributions to  
1137 general geology, 1914. USGS Professional Paper 90-G.

1138

1139 Stebinger, E. 1917a. Contributions to economic geology, 1916, Part II, Mineral fuels-- Anticlines  
1140 in the Blackfoot Indian Reservation, Montana. USGS Bulletin 641-J.

1141

1142 Stebinger, E. 1917b. Contributions to economic geology, 1916, Part II, Mineral fuels.  
1143 Possibilities of oil and gas in north-central Montana. USGS Bulletin 641-C.

1144

1145 Tokarsky, O. 1974. Hydrogeology of the Lethbridge-Fernie area, Alberta. Alberta Research.

1146

1147 Toth, J., and Corbet, T. 1986. Post-Paleocene evolution of regional groundwater flow-systems  
1148 and their relation to petroleum accumulations, Taber area, southern Alberta, Canada. Bulletin of  
1149 Canadian Petroleum Geology **34**:339–363.

1150

1151 Tovell, W.M. 1956. Some Aspects of the Geology of the Milk River and Pakowki Formations  
1152 (Southern Alberta). Thesis, University of Toronto.

1153



- 1154 Tuck, L.K. 1993. Reconnaissance of geology and water resources along the north flank of the  
1155 Sweet Grass Hills, north-central Montana. USGS Water-Resources Investigations Report: 93-  
1156 4026.  
1157
- 1158 Weed, W.H. 1899. Fort Benton folio, Montana. Report GF-55. USGS Folios of the Geologic  
1159 Atlas: 55.  
1160
- 1161 WEBLEX Canada 2013. Natural Resources Canada. Government of Canada.  
1162 <http://weblex.nrcan.gc.ca>.  
1163
- 1164 Williams, M.Y., and Dyer, W.S. 1930. Geology of southern Alberta and southwestern  
1165 Saskatchewan. Memoir 163, Geological Survey of Canada.  
1166
- 1167 Zimmerman, E. 1967. Water resources of the Cut Bank area, Glacier and Toole Counties,  
1168 Montana. Montana Bureau of Mines and Geology, Bulletin 60.